

2017-05-12

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Springer Nature Switzerland AG.

<https://doi.org/10.1007/s13201-017-0569-6>

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Origin and mechanisms of high salinity in Hombolo Dam and groundwater in Dodoma municipality Tanzania, revealed

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Received: 29 June 2015 / Accepted: 24 April 2017 / Published online: 12 May 2017
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Abstract The Hombolo dam (HD), in central Tanzania, is a shallow reservoir characterized by high salinity that limits its use for human activities. The origin of the salinity, mechanisms of reaching and concentrating in the dam remain unclear. These were assessed using hydrogeochemical facies, water type evolutions and mapping. The source of HD salinity was identified to be shallow groundwater (SG) and runoff from a seasonal floodplain with NaCl-rich lithological materials, along Little Kinyasungwe River that feeds the dam. The NaCl-rich lithological units, about 5–7 km upstream of the dam, were highly concentrated with NaCl to the extent that the local community was commercially separating table salt from them. The physicochemical parameters from these NaCl-rich lithological materials were well represented in HD and nearby groundwater sources, which suggests active water interactions. Water type evolution and surface hydrology assessments clearly showed that SG in the salty-floodplain was influenced by evaporation (ET) and was periodically carried to the HD. Clearly; HD water had high chemical similarity with the nearby SG. This agrees with previous studies that HD is partly fed by the local aquifer. However, this is the first

attempt at mapping its physical origin. The origin of HD salinity was further supported by the spatial distribution of electrical conductivity (EC), where very high EC (up to 21,230 μScm^{-1}) was recorded in SG within the NaCl-rich lithological unit while water sources far away from the NaCl-rich materials had much lower EC values. Thus, the study disagrees with previous conclusions that HD salinity was solely due to high dam surface ET but is primarily due to geological reasons. Comparisons of HD with a nearby Matumbulu dam (MD), another earthen dam in climatologically similar settings, reveals that MD water was less saline/mineralised. This further shows that HD high salinity is most likely a geologic phenomenon, but local climatic factors, namely high ET, decreasing rainfall and warming trends are likely to have concentrated the salts further. Although HD is widely/ideally used for grape vine irrigation, it was clearly revealed that its prolonged usage would potentially affect the soil and grape productivity due to high salinity.

Keywords · Hombolo dam · salinity · groundwater · origin · livelihood

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Introduction

Water is a universal solvent in which most matters dissolve (Kumar 2013; Parmar 2013). Thus, its chemical compositions depend on environmental and aseptic conditions and processes that potentially modify it as it moves along the catchment (Biswas 2004; Kumar 2013). Thus, the best way to manage water quality is tracing its catchment movements (Parmar 2013). In recent years, however, global freshwater quality has been worsened by humans and ongoing climate change impacts (Taylor et al. 2012, Shemsanga et al. 2015). Yet, even without these problems,

water quality/salinity management; is difficult as water molecules constantly move between different states and locations (Biswas 2004 2008).

In arid and semi-arid zones, water quality limits water supply options and uses (Prasanna et al. 2011). Thus, management of such valuable water resources calls for an understanding of its hydro-geochemistry and evolutions (Prasanna et al. 2011). Since different water uses/needs require specific quality, the origin of freshwater quality is of global interest (Ghalibaf and Moussavi 2014). For instance, while high water salinity affects soil/crop productivity, hard water causes scum, curd, yellowing of fabrics, scales in heaters/boilers and pipes. It also toughens vegetables and does not easily form lather with soap. Yet, too soft water is often corrosive (Mulwa et al. 2013). Thus, assessing the origin of water quality and allocating its suitable uses is crucial (Mulwa et al. 2013).

Although Hombolo dam (HD) high salinity has been debated for over 20 years, its origin remains unclear. While some studies associated it with high dam surface evaporation (ET) (Shindo 1991), others linked it to geological origin, but would not clearly indicate the mechanisms (Nkotagu 1996). All, previous studies, however, did not show the spatial source of the salts that would end up being concentrated by ET nor assess water suitability for various uses. Furthermore, available data at Hombolo Irrigation Scheme (HIS), which started in 1995, do not suggest a significant rainfall dilution and that the dam retains high salinity year-round (Shindo 1991; Shemsanga et al. 2015). This suggests that the water feeding the dam could already be saline, details of which remains to be investigated.

Yet, irrigation officers and irrigators reported unexplained progressive increases in HD salinity recently. This has reduced seasonality for which HD is used for irrigation (mostly grape). Village leaders, elders and fishermen argue of reduced HD fish harvests recently while well owners around HD report increased pumpage of saltier water recently (Personal communication with Mr. Ntuza of HIS). Yet, HD is a vital discharge site of Makutupora aquifer, the main aquifer supplying water to Dodoma municipality (DM) (Shindo 1991). Thus, for the integrity of the city, the origin of salts, their evolutions, movements and mechanisms of concentrating in HD and groundwater must be understood. Nevertheless, these remain unclear which impedes sustainable water resources management.

Despite the importance of HD, no current study has assessed sources of its high salinity, transportation and concentrating mechanisms in the dam. This study, therefore, intended to critically investigate and map the origin and mechanisms of high salinity in the HD catchment so as to help the local populace to support their major livelihood support systems that mostly depend on it.

Case study description

This study was carried in HD catchment within Dodoma Municipality, which is located between 775,000–945,000 and 9,280,000–9,380,000 (UTM, Arc 1960) (Fig. 1). Climatologically, it is a semi-arid region with low uni-modal rains (550 mm/yr.) falling between October and May. Mean daily temperature is lowest in July (about 13.0 °C) and highest in November (about 30.6 °C). ET rate averages about 2000 mm yr^{-1} ; roughly 4 times annual rainfall (Shemsanga et al. 2015). Except for HD, the area lacks permanent surface water bodies and the river network is largely seasonal flowing only during wet seasons and a few weeks thereafter. Thus, human activities are heavily reliant on HD and groundwater. Geologically, it occurs in the fractured, faulted and tectonically active crystalline basement of the Dodoma craton. The meta-sedimentary rocks are mainly quartzites, ironstones, mica-ceous quartzites, quartzo-feldspathic schists and ferruginous quartzites (Nkotagu 1996, Shemsanga et al., 2015).

Methodology

Twenty-seven (27) water samples from boreholes (BHs), HD, Matumbulu dam (MD), shallow dug wells (SDWs) and the major tributaries leading to Little Kinyasungwe River (LKR) and which flows to HD; were collected in March, 2014. HD was also sampled during the peak of the dry season, i.e. September, 2014 to enable useful water quality parameters temporal comparisons. Water quality comparisons between HD and MD, another earthen dam in similar environmental settings (about 38 km apart), were also done. Scichem STC multi-parameter tester kit was used for in situ measurements of pH, total dissolved solids (TDS) and electrical conductivity (EC). Sampling sites were marked by handheld GPS receiver, Garmin eTrex 20 (UTM, Arc 1960). Water samples were analyzed at the Ministry of Water Laboratory in Dodoma following standard methods described by Elisante and Muzuka (2016). Laboratory results were checked for consistency via electrical balance error (EBE) (Eq. 1) and were required to be <10 (Kaka et al. 2011). Surfer V.11[®] was used to extrapolate EC and pH spatial patterns using kriging tool (Fig. 2a, b). Water suitability for irrigation was assessed using standard indices (Eqs. 2–5), all concentrations (conc.) expressed in meq/L. Domestic water was assessed based on national (URT 2007) and World Health Organization (WHO 2008) permissible limits (PLs) and maximum permissible limits (MPLs) (Table 1). Since local water sources are mainly used for grape irrigation, and because grapes are sensitive to high salinity, chloride and sodium toxicity; their suitability was assessed using standard methods described by Goldammer (2015). Grapes require Na^+ and Cl^- in irrigation water

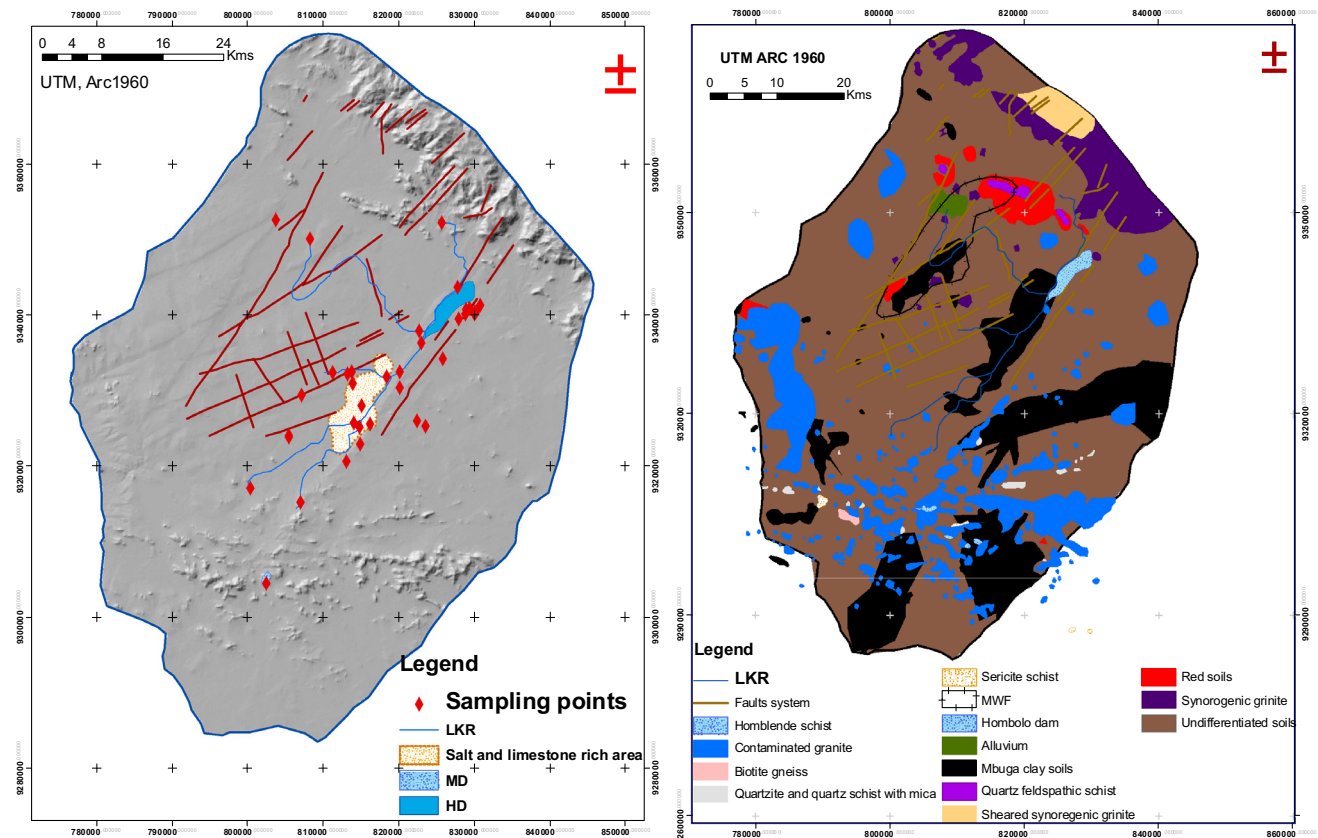


Fig. 1 A map showing the location of the study area, sampling points, geological map and Hombolo and Matumbulu dams of Dodoma Municipality—Tanzania

be to <20.01 and <3.95 meq/L, respectively, while pH should range from 6 to 8.5. Generally, EC and Cl⁻ of (0–2000 and <3.95) (20,001–30,000 and 3.95– <14.95), and (>3000 and >14.95) (μScm⁻¹ and meq/L), cause, no, medium and severe problems for grape vine productivity, respectively (Goldammer 2015).

$$\%EBE = \frac{\left(\sum \text{Cations} - \sum \text{Anions}\right)}{\sum \left(\text{Cations} + \sum \text{Anions}\right)} (\text{meq/L}) \times 100 \quad (1)$$

Graphical compositions of cations and anions in water were assessed using a diamond-shaped PIPER diagram embedded in AquaChem V.4.0[®]. This was also used to classify “hydrochemical facies” or “water types” based on the dominant ions (Piper 1944). Gibb’s and Chadha’s diagrams were used to assess major water evolutions/mechanisms, relationships between water chemistry and aquifer lithology, and to differentiate contributions of rock–water interactions, ET and precipitation dominances (Gibbs 1970; Chadha 1999). Ion exchange and reverse ion exchange reactions were assessed using Chloro-alkaline indices 1 and 2 (CAI 1 and CAI 2, Eqs. 6 and 7), all conc.

| Indices | Formulae | References |
|---------------------------------|---|---|
| Sodium absorption ratio (SAR) | $SAR = \frac{Na}{\sqrt{(Ca+Mg)/2}} \quad (2)$ | Inayathulla and Paul (2013); Sadashivaiah et al. (2008) |
| Residual sodium carbonate (RSC) | $RCS = (HCO_3 + CO_3) - (Ca + Mg) \quad (3)$ | Inayathulla and Paul (2013); Kaur and Singh (2011) |
| Percent sodium (%Na) | $\% Na = \left[\frac{Na}{Ca+Mg+Na+K} \right] \times 100 \quad (4)$ | Sadashivaiah et al. (2008); Kaur and Singh (2011) |
| Kelly’s ratio (KR) | $KR = \frac{Na}{(Ca+Mg)} \quad (5)$ | Inayathulla and Paul (2013); Kaur and Singh (2011) |

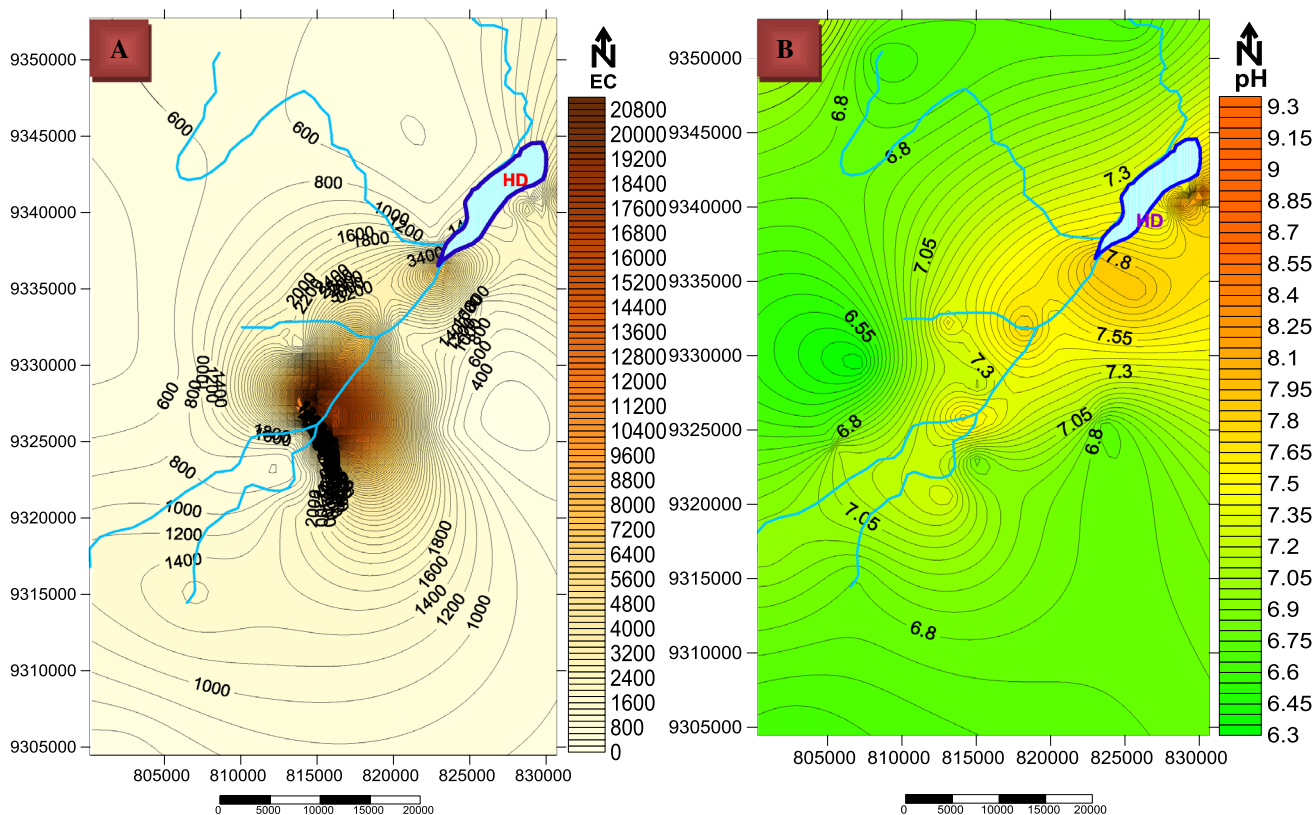


Fig. 2 a Spatial EC distribution, b Spatial pH distribution from the study catchment

expressed in meq/L. Field observations (FOs) and Focus group discussions (FGDs) were held for five (5) months covering both the dry season (September–October 2013) and wet season (December, 2013 - February, 2014). When the use of local Gogo dialect was necessary, due to communication barriers, the field assistant (Mr. Malima) provided translations. Human and geological processes and activities upstream of HD were assessed by analyzing and documenting all potential sources of HD salinity. An extensive literature review was also undertaken.

$$\text{CAI 1} = \text{Cl} - \text{Na} + \text{K/Cl} \quad (6)$$

$$\text{CAI 2} = \text{Cl} - \text{Na} + \text{K/SO}_4 + \text{HCO}_3 + \text{CO}_3 + \text{NO}_3 \quad (7)$$

Total hardness (TH) is vital for assessing suitability of water sources for domestic, agricultural and industrial uses and it was assessed using Eq. 8, all concentrations were expressed in ppm. Generally, water sources with TH values of (0–75), (>75–150) (>150–300) (>300) were classified as soft, moderately hard, hard, and very hard, respectively (Kaka et al. 2011).

$$\text{TH} = 2.497 \times [\text{Ca}^{2+}] + 4.115 \times [\text{Mg}^{2+}] \quad (8)$$

Results

Physical and hydrogeochemical parameters

Table 1 summarizes water samples analytical results (range, mean, median and SD values) for major physico-chemical water quality parameters. The mean cations and anions followed ($\text{Na}^+ > \text{Mg}^{2+} > \text{Ca}^{2+} > \text{K}^+$) and ($\text{Cl}^- > \text{HCO}_3^- > \text{SO}_4^{2-} > \text{NO}_3^-$) order and constituted (14.34, 5.19, 4.69, 0.40) and (11.83, 6.33, 3.24, 1.13) meq/L, respectively (Table 1). Mean pH, EC, TDS, TH, SAR, %Na, RSC and KR were 7.4, 3127.8 μScm^{-1} , 2004.5, 533.1 ppm, 6.6, 51%, –3.5 and 1.7, respectively (Tables 1, 2). The highest total cations (TC) and total anions (TA) occur in the NaCl-rich lithological unit where table salt is mined (Table 1, Plate 1). HD water quality varied with seasons where pH and EC were slightly lower during the wet season than in the dry season, (8.0 and 8.5) and (3298 and 4006 μScm^{-1}), respectively (Table 1). The middle of the catchment had highest EC (Fig. 2a).

From Table 1, SD for Na^+ and K^+ were 15.19 and 0.46 and ranged from 2.04 to 65.10 and 0.06 to 2.27 meq/L, respectively. SD of Ca^{2+} and Mg^{2+} were 2.81 and 3.85 and

Table 1 Field and analytical results of the key physicochemical water quality parameters

| SN | Sample ID | EC μS _{cm} ⁻¹ | EC Class | pH | Na ⁺ (meq) | Ca ²⁺ (meq) | Mg ²⁺ (meq) | K ⁺ (meq) | Cl ⁻ (meq) | SO ₄ ²⁻ (meq) | HCO ⁻ (meq) | NO ₃ ⁻ (meq) | CO ₃ ²⁻ (meq) | TDS (ppm) | TH (ppm) | TA (meq) | TC (meq) | TA-TC (meq) |
|--------|--------------------|--------------------------------------|-------------|------|--------------------------|---------------------------|---------------------------|-------------------------|--------------------------|--|---------------------------|---------------------------------------|--|--------------|-------------|----------|----------|----------------|
| 1 | HD-dry | 4006 | C4 | 8.5 | 20.23 | 2.15 | 2.86 | 0.56 | 14.28 | 5.40 | 4.70 | 0.08 | 0.80 | 2566 | 250.82 | 25.8 | 25.26 | 0.54 |
| 2 | HD-wet | 3298 | C4 | 8.0 | 17.75 | 2.68 | 4.25 | 0.48 | 13.99 | 3.86 | 5.60 | 0.04 | 0.00 | 2112 | 346.83 | 25.16 | 23.49 | 1.67 |
| 3 | MD | 209 | C1 | 6.6 | 10.55 | 4.53 | 4.79 | 0.41 | 9.60 | 2.48 | 5.96 | 0.72 | 0.00 | 134 | 465.91 | 20.28 | 18.76 | 1.52 |
| 4 | Paul Kijuu (BH) | 767 | C3 | 6.3 | 2.07 | 3.05 | 2.56 | 0.30 | 2.13 | 0.37 | 4.33 | 0.56 | 0.00 | 491 | 280.79 | 7.98 | 7.39 | 0.59 |
| 5 | M/Nyika vill. (BH) | 1867 | C3 | 7.5 | 12.09 | 3.59 | 7.08 | 0.45 | 8.80 | 4.58 | 8.02 | 0.49 | 0.00 | 1197 | 533.67 | 23.21 | 21.89 | 1.32 |
| 6 | M/Nyika (SW) | 2476 | C4 | 7.5 | 16.25 | 3.79 | 11.44 | 1.51 | 14.22 | 5.79 | 10.36 | 0.62 | 0.00 | 1588 | 761.76 | 32.99 | 30.99 | 2 |
| 7 | Nzuguni Pol (BH) | 1728 | C3 | 7.3 | 18.80 | 5.69 | 4.67 | 0.37 | 11.64 | 4.88 | 9.48 | 2.42 | 0.00 | 1112 | 518.64 | 29.53 | 28.42 | 1.11 |
| 8 | M/Nyika DW Wowa | 1140 | C3 | 7.4 | 16.30 | 5.98 | 9.67 | 2.27 | 10.35 | 9.40 | 10.62 | 0.39 | 0.00 | 729 | 782.65 | 34.22 | 30.76 | 3.46 |
| 9 | Nzuguni Vil (BH) | 1630 | C3 | 6.8 | 4.25 | 2.83 | 1.22 | 0.19 | 2.80 | 0.73 | 3.76 | 0.62 | 0.00 | 1045 | 202.48 | 8.49 | 7.91 | 0.58 |
| 10 | Kitelela I (BH) | 2534 | C4 | 7.5 | 16.95 | 3.98 | 12.45 | 0.25 | 21.69 | 2.47 | 5.23 | 1.27 | 0.00 | 1622 | 821.86 | 33.63 | 30.66 | 2.97 |
| 11 | Kitelela II (BH) | 2171 | C3 | 7.3 | 13.61 | 3.92 | 12.32 | 0.23 | 19.71 | 2.29 | 5.22 | 1.22 | 0.00 | 1394 | 812.28 | 30.08 | 28.44 | 1.64 |
| 12 | Happy Hom (BH) | 894 | C3 | 7.3 | 9.29 | 6.14 | 3.46 | 0.39 | 8.21 | 1.79 | 5.87 | 0.90 | 0.00 | 572 | 479.96 | 19.28 | 16.77 | 2.51 |
| 13 | LGT 1 (BH) | 3261 | C4 | 7.6 | 11.14 | 6.55 | 1.04 | 0.19 | 11.23 | 2.17 | 4.58 | 0.28 | 0.00 | 2093 | 379.66 | 18.92 | 18.26 | 0.66 |
| 14 | LGT 2 (BH) | 2787 | C4 | 7.8 | 8.61 | 5.79 | 5.05 | 0.06 | 9.44 | 1.04 | 6.21 | 0.20 | 0.00 | 1780 | 542.31 | 19.51 | 16.89 | 2.62 |
| 15 | LGT 3 (BH) | 2284 | C4 | 9.3 | 3.20 | 1.30 | 2.37 | 0.09 | 3.41 | 1.02 | 1.40 | 0.27 | 0.80 | 1461 | 183.68 | 6.96 | 6.9 | 0.06 |
| 16 | Kadirimsimba (BH) | 2289 | C4 | 7.8 | 9.99 | 6.19 | 1.24 | 0.39 | 5.51 | 1.38 | 6.44 | 1.54 | 0.00 | 1466 | 371.68 | 17.81 | 14.87 | 2.94 |
| 17 | Maria Clara (BH) | 1507 | C3 | 7.4 | 6.21 | 3.37 | 3.10 | 0.17 | 5.40 | 1.44 | 4.35 | 0.08 | 0.00 | 964 | 323.68 | 12.85 | 11.27 | 1.58 |
| 18 | Hombolo Agr. (BH) | 2089 | C3 | 7.9 | 26.65 | 1.22 | 1.09 | 0.06 | 18.56 | 1.14 | 5.35 | 0.04 | 0.00 | 1342 | 115.91 | 29.02 | 25.09 | 3.93 |
| 19 | Kitelela (SDW) | 1767 | C3 | 7.3 | 4.89 | 9.19 | 10.40 | 0.41 | 11.63 | 2.74 | 8.91 | 1.03 | 0.00 | 1132 | 979.83 | 24.89 | 24.31 | 0.58 |
| 20 | M/Makulu (BH) | 1657 | C3 | 7.1 | 17.06 | 14.97 | 6.99 | 0.20 | 25.98 | 1.98 | 7.17 | 2.16 | 0.00 | 1061 | 1098.87 | 39.22 | 37.29 | 1.93 |
| 21 | M/Makulu (SDW) | 817 | C3 | 6.7 | 5.35 | 4.05 | 2.58 | 0.13 | 2.79 | 1.09 | 7.48 | 0.70 | 0.00 | 523 | 331.97 | 12.11 | 12.06 | 0.05 |
| 22 | Ipala TAG (BH) | 1006 | C3 | 7.9 | 6.06 | 2.78 | 1.49 | 0.12 | 2.96 | 1.71 | 4.51 | 1.27 | 0.00 | 644 | 213.56 | 10.45 | 10.45 | 0 |
| 23 | Msisi Spring | 668 | C2 | 7.3 | 2.28 | 3.32 | 2.38 | 0.31 | 1.44 | 1.46 | 4.67 | 0.00 | 0.00 | 427 | 285.54 | 8.29 | 7.57 | 0.72 |
| 24 | Laswai (BH) | 812 | C3 | 6.8 | 2.08 | 3.03 | 2.55 | 0.28 | 2.13 | 0.37 | 4.35 | 0.56 | 0.00 | 523 | 279.38 | 7.94 | 7.41 | 0.53 |
| 25 | Mwala (BH) | 799 | C3 | 7.1 | 2.04 | 3.04 | 2.63 | 0.31 | 2.11 | 0.36 | 4.36 | 0.58 | 0.00 | 511 | 283.99 | 8.02 | 7.41 | 0.61 |
| 26 | Dug well at SPP I | 21,230 | C4 | 7.2 | 58.50 | 8.03 | 9.71 | 0.36 | 40.05 | 14.04 | 8.30 | 5.23 | 0.00 | 13,621 | 887.49 | 76.6 | 67.62 | 8.98 |
| 27 | Dug well at SPP II | 18,756 | C4 | 7.4 | 65.10 | 5.44 | 10.86 | 0.37 | 39.30 | 11.40 | 13.75 | 7.37 | 0.00 | 12,012 | 1858.67 | 81.77 | 71.82 | 9.95 |
| Aver. | | 3127.74 | C4 | 7.43 | 14.34 | 4.69 | 5.19 | 0.40 | 11.83 | 3.24 | 6.33 | 1.13 | 0.10 | 2004.52 | 533.11 | 24.63 | 22.59 | 2.04 |
| Min. | | 209.00 | | 6.30 | 2.04 | 1.22 | 1.04 | 0.06 | 1.44 | 0.36 | 1.40 | 0.00 | 0.00 | 134.00 | 115.91 | 6.96 | 6.9 | 0 |
| Maxi. | | 21,230.0 | | 9.30 | 65.10 | 14.97 | 12.45 | 2.27 | 40.05 | 14.04 | 13.75 | 7.37 | 0.90 | 13,621.0 | 1858.67 | 81.77 | 71.82 | 9.95 |
| Median | | 1767.00 | | 7.40 | 10.55 | 3.92 | 3.46 | 0.31 | 9.60 | 1.98 | 5.60 | 0.62 | 0.00 | 3178.81 | 379.66 | 20.28 | 18.76 | 1.52 |
| SD | | 4958.28 | | 0.60 | 15.19 | 2.81 | 3.85 | 0.46 | 10.36 | 3.43 | 2.61 | 1.64 | 0.27 | 1132.00 | 376.89 | 18.42 | 16.27 | 2.39 |

Table 1 continued

| SN | Sample ID | EC μScm^{-1} | EC Class | pH | Na^+ (meq) | Ca^{2+} (meq) | Mg^{2+} (meq) | K^+ (meq) | Cl^- (meq) | SO_4^{2-} (meq) | HCO_3^- (meq) | NO_3^- (meq) | CO_3^{2-} (meq) | TDS (ppm) | TH (ppm) | TA (meq) | TC (meq) | TA-TC (meq) | |
|----|---|----------------------------|-----------------------------------|---------|------------------------|---------------------------|---------------------------|-----------------------|------------------------|-----------------------------|---------------------------|--------------------------|-----------------------------|--------------|-------------|----------|----------|----------------|--|
| | National standard | n.m | | 6.5–9.2 | 26.10 | 14.97 | 8.23 | 0.51 | 22.57 | 16.67 | n.m | 1.21 | n.m | 2000 | <75, soft | 25.8 | 25.26 | 0.54 | |
| | WHO standard | n.m | | 6.5–8.5 | 8.70 | 9.98 | 12.34 | 0.31 | 16.93 | 8.33 | n.m | 0.81 | n.m | 500 | | 25.16 | 23.49 | 1.67 | |
| | Number of samples, classes and percentages | | | | | | | | | | | | | | | | | | |
| | | | 100–250, C1 (3.7%) (Excellent) | | | | | | | | | | | | | | | | |
| | | | 251–750, C2, (3.7%) (Good) | | | | | | | | | | | | | | | | |
| | | | 751–2250, C3 (55.6%) (Average) | | | | | | | | | | | | | | | | |
| | | | >2251, C4 37% (Bad) | | | | | | | | | | | | | | | | |

n.m not mentioned, BH borehole, SDW shallow dug well, SPP salt processing plant

ranged from (1.22 to 14.97) and 1.04 to 12.45 meq/L, respectively (Table 1). SO_4^{2-} ranged from 0.36 to 14.04 and had SD of 3.43 meq/L (Table 1). Cl^- ranged from 1.44 to 40.05 and had SD of 10.36 while NO_3^- ranged from 0.0 to 7.37 and had SD of 1.64 meq/L, respectively (Table 1). The highest NO_3^- and Cl^- were recorded at the two SDWs within the limestone and salt materials and processing activities (LSPMA) (Table 1). Mostly nil to traces of CO_3^{2-} (0–0.8 meq/L) were recorded in most waters while HCO_3^- was the dominant anion varying from 1.4 to 13.75 with SD of 2.61 meq/L (Table 1).

Table 1 shows that 29.6% of samples were above the national and WHO PL for NO_3^- . While all samples were within national PL for SO_4^{2-} , 11.1% were above the WHO PL. With regards to Cl^- , 22.2 and 11.1% of the samples were above the PL for WHO and national standards, respectively. In terms of cations, 3.7% of the samples were above Ca^{2+} PL for both WHO and national PL while 3.7 and 22.2% were above the WHO and national PL for Mg^{2+} , respectively. Furthermore, 48.1 and 11.1% of samples were above WHO and national PL for K^+ , respectively. 72.7 and 11.1% of samples were above the WHO and national PL for Na^+ , respectively which represent a big difference between the two guidelines. Finally, 66.7, 29.3, and 3.7% of samples had very hard; hard and moderately hard TH values, respectively and had a mean value of 533.1 and a range of 115.91–1858.7 (Table 1). This survey indicated that prior to this study no water quality assessment existed for about 63% of the samples and that local people mostly used water of unknown quality (Field observations s and FGDs).

Major mechanisms controlling water quality, evolution and hydrogeochemical facies

Figure 4 shows that most water samples were in rock dominance. However, for both wet and dry seasons, HD and samples collected from shallow groundwater (SG) in the LSPMA were in the ET dominance. Unlike HD, MD spanned in the rock dominance, and closer to precipitation and had the least TDS abundance. Notice that no samples were in the precipitation dominance and that most samples had high TDS and occurred in the upper half of Gibbs plots (Fig. 4). The order of water chemical mechanism dominance was seawater ET, reverse ion exchange, recharge water and to a limited extent, ion exchange, respectively (Fig. 3b). Figure 7 shows that the dominant cations and anions were ($\text{Na}^+ + \text{K}^+$) and Cl^- , respectively and the major water type was Na–Cl followed by mixed CaMgCl and CaHCO₃ in that order. For both seasons, HD and SGs from the LSPMA had Na–Cl water type while MD was represented by CaHCO₃ and seasonal

Table 2 Major statistics for SAR, %Na, RSC, KR and Indices

| SN | %EBE | SAR | SAR Class | %Na | %Na Class | RSC | RSC Class | KR | Na/Cl | UTM-X | UTM-Y |
|--|-------|---|-----------|--|-----------|--|-----------|--------|--------|---------|-----------|
| 1 | 1.051 | 12.77 | S2 | 78.404 | C4 | 0.491 | C1 | 4.034 | 9.404 | 828,810 | 9,340,236 |
| 2 | 1.575 | 9.53 | S1 | 70.532 | C4 | -0.430 | C1 | 2.560 | 6.624 | 828,810 | 9,340,236 |
| 3 | 3.459 | 5.39 | S1 | 48.818 | C3 | -3.251 | C1 | 1.133 | 2.331 | 802,425 | 9,304,467 |
| 4 | 3.920 | 1.24 | S1 | 25.930 | C2 | -1.285 | C1 | 0.369 | 0.678 | 807,153 | 9,329,447 |
| 5 | 2.936 | 5.24 | S1 | 52.098 | C3 | -2.648 | C1 | 1.133 | 3.366 | 820,097 | 9,332,557 |
| 6 | 3.137 | 5.89 | S1 | 49.259 | C3 | -4.872 | C1 | 1.067 | 4.285 | 820,134 | 9,330,485 |
| 7 | 1.919 | 8.26 | S1 | 63.637 | C4 | -0.884 | C1 | 1.813 | 3.301 | 813,122 | 9,320,581 |
| 8 | 5.307 | 5.83 | S1 | 47.640 | C3 | -5.025 | C1 | 1.042 | 2.726 | 820,134 | 9,330,485 |
| 9 | 3.585 | 2.98 | S1 | 50.024 | C3 | -0.291 | C1 | 1.049 | 1.500 | 814,901 | 9,322,943 |
| 10 | 4.609 | 5.91 | S1 | 50.410 | C3 | -11.195 | C1 | 1.032 | 4.257 | 813,190 | 9,332,206 |
| 11 | 2.784 | 4.78 | S1 | 45.253 | C3 | -11.013 | C1 | 0.838 | 3.470 | 813,751 | 9,332,475 |
| 12 | 6.914 | 4.24 | S1 | 48.202 | C3 | -3.723 | C1 | 0.968 | 1.513 | 829,294 | 9,340,903 |
| 13 | 1.271 | 5.72 | S1 | 58.863 | C3 | -3.009 | C1 | 1.468 | 1.701 | 830,129 | 9,340,448 |
| 14 | 7.179 | 3.70 | S1 | 44.141 | C3 | -4.628 | C1 | 0.795 | 1.488 | 830,061 | 9,340,445 |
| 15 | 0.385 | 2.36 | S1 | 45.929 | C3 | -1.469 | C1 | 0.871 | 2.455 | 830,083 | 9,340,536 |
| 16 | 9.019 | 5.19 | S1 | 56.095 | C3 | -0.987 | C1 | 1.345 | 1.615 | 828,950 | 9,340,759 |
| 17 | 6.577 | 3.45 | S1 | 48.328 | C3 | -2.122 | C1 | 0.960 | 1.844 | 827,873 | 9,339,432 |
| 18 | 7.272 | 24.76 | S3 | 91.815 | C5 | 3.031 | C3 | 11.502 | 21.797 | 830,014 | 9,340,236 |
| 19 | 1.183 | 1.56 | S1 | 19.660 | C1 | -10.673 | C1 | 0.250 | 0.533 | 813,988 | 9,330,999 |
| 20 | 2.532 | 5.15 | S1 | 43.492 | C3 | -14.796 | C1 | 0.777 | 1.140 | 822,486 | 9,326,018 |
| 21 | 0.260 | 2.93 | S1 | 44.157 | C3 | 0.842 | C1 | 0.806 | 1.319 | 823,498 | 9,325,344 |
| 22 | 1.051 | 4.15 | S1 | 58.031 | C3 | 0.245 | C1 | 1.421 | 2.182 | 825,744 | 9,334,181 |
| 23 | 4.599 | 1.35 | S1 | 27.479 | C2 | -1.035 | C1 | 0.400 | 0.687 | 803,741 | 9,352,685 |
| 24 | 3.474 | 1.25 | S1 | 26.210 | C2 | -1.237 | C1 | 0.373 | 0.687 | 805,422 | 9,323,988 |
| 25 | 3.934 | 1.21 | S1 | 25.426 | C2 | -1.314 | C1 | 0.359 | 0.670 | 805,475 | 9,323,937 |
| 26 | 6.23 | 19.64 | S3 | 76.371 | C4 | -9.440 | C1 | 3.298 | 7.285 | 815,077 | 9,327,954 |
| 27 | 6.48 | 22.80 | S3 | 79.614 | C4 | -2.550 | C1 | 3.994 | 11.967 | 816,132 | 9,325,644 |
| Aver. | 3.60 | 6.57 | S1 | 50.96 | C3 | -3.45 | C1 | 1.69 | 3.73 | | |
| Min. | | 1.21 | | 19.66 | | -14.80 | | 0.25 | 0.53 | | |
| Max. | | 24.76 | | 91.82 | | 3.03 | | 11.50 | 21.80 | | |
| Median | | 5.15 | | 48.82 | | -2.12 | | 1.04 | 2.18 | | |
| SD. | 2.61 | 6.32 | | 17.80 | | 4.34 | | 2.21 | 4.57 | | |
| Number of samples, classes and percentages | | <10 = (S1, Excellent), 23 samples, 85.2% | | C1 Excellent (< 20), 1 samples, 3.7% | | <1.25 (C1, Good safe), 26 sample, 96.3% | | | | | |
| | | 11–12 = S2 (Good), 1 samples, 3.7% | | C2 Good (20–40), 4 samples, 7.4% | | 1.25–2.5 (C2, Medium) 0 sample 0% | | | | | |
| | | 19–26 = S3 (Average), 3 sample, 11.1% | | C3 Permissible (40–60), 16 samples, 59.3% | | >2.5 (C3, Unsuitable), 1 samples, 3.7% | | | | | |
| | | >27 = S4(Bad), 0 samples, 0% | | C4 Doubtful (60–80), 5 samples, 7.7% | | | | | | | |
| | | | | C5 Unsuitable (>80), 1 samples, 3.7% | | | | | | | |
| References | | Sadashivaiah et al. (2008) | | Inayathulla and Paul (2013); Sadashivaiah et al. (2008) | | Kaur and Singh (2011); Sadashivaiah et al. (2008) | | | | | |

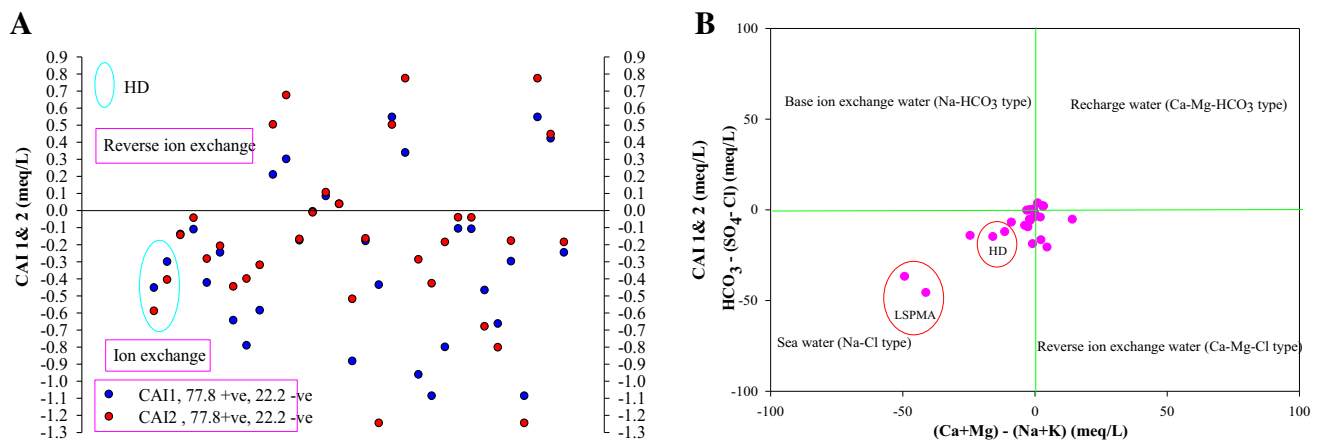


Fig. 3 a CAI 1 and CAI 2, b Chadha diagrams for hydrogeochemical facies and evolutions

dilution did not change the general chemistry (Fig. 7a). While one sample had mixed CaNaHCO_3 , none had CaCl and NaHCO_3 types.

Figure 5a shows that the molar ratio between Na/Cl were mostly >1 and plotted below the (equiline 1:1). Figure 6d, p, j show that most samples span above equilines of HCO_3^- vs. $(\text{Na}^+ + \text{K}^+)$, $(\text{Ca}^{2+} + \text{Mg}^{2+})$ vs. $(\text{HCO}_3^- + \text{SO}_4^{2-})$ and $(\text{Ca}^{2+} + \text{Mg}^{2+})$ vs. TC, respectively. Further, most samples were below and above $(\text{HCO}_3^- + \text{SO}_4^{2-})$ vs. $(\text{Na}^+ + \text{K}^+)$ and $(\text{Ca}^{2+} + \text{Mg}^{2+})$ vs. HCO_3^- equilines, respectively. Notice the strong $-ve$ correlation ($r^2 = 0.73$) between $(\text{Ca}^{2+} + \text{Mg}^{2+}) - (\text{HCO}_3^- + \text{SO}_4^{2-})$ vs. $(\text{Na}^+ \text{ to } \text{Cl}^-)$ (Fig. 6f) and high $+ve$ correlations between $(\text{Na}^+ + \text{K}^+)$ vs. $(\text{Cl}^- + \text{SO}_4^{2-})$, $(\text{Na}^+ + \text{K}^+)$ vs. TC and $(\text{Ca}^{2+} + \text{Mg}^{2+})$ vs. TC of (Fig. 6i, $r^2 = 0.9$), (Fig. 6m, $r^2 = 0.93$) and (Fig. 6j, $r^2 = 0.9$), respectively. Notice also the $+ve$ correlation ($r^2 = 0.56$) between Cl^- and NO_3^- and that the highest values were recorded at the SG within the LSPMA (Table 1). A plot of Cl^- vs. TA and $(\text{HCO}_3^- + \text{SO}_4^{2-})$ vs. TA revealed that all samples plotted below the equiline, Fig. 5r, t, respectively. Further, 81.5% of samples in a graph of $(\text{Ca}^{2+} + \text{Mg}^{2+})$ vs. Cl^- plot above the equiline and that $(\text{Ca}^{2+} + \text{Mg}^{2+})$ increases with salinity (Fig. 5b). Further, Na/Cl slightly decreases with Cl^- (Fig. 6s). The Na/Ca molar ratio of about 52% of the samples locally were >2 with mean and range of 3.73 and 0.53–21.8, respectively (Table 2). Figure 3a shows that most samples (including HD) had $-ve$ CAI 1 and CAI 2 values (70.4%) while a few had $+ve$ values (29.6%) and were located in the ion exchange and reverse ion exchange, respectively. Finally, Fig. 7b shows the Schoeller diagram which, irrespective of various concentrations from different samples, similar patterns could be traced.

Irrigation water suitability indices

Table 2 shows that from different indices, local water suitability for irrigation is highly variable. While water with KR of <1 is suitable for irrigation, current results show that 55.6% of samples had KR of >1 with HD and samples from LSPMA having KR of >2 and >3 , respectively. Based on the SAR index, 85.2% of samples were grouped as S1 while 3.7 and 11.1% had S2 and S3, respectively. By following %Na index, 3.7, 7.4, 59.3, 7.7 and 3.7% of the samples were classified as C1, C2, C3, C4 and C5, respectively (Table 2). Based on RSC classification, 96.3% of the samples were good for irrigation whilst 3.7% were bad/unusable (Table 2). Notice that based on %Na and RSC indices, HD water was suitable and unsuitable for irrigation, respectively (Table 2). From most indices, SN 18 and HD were the most unfit for irrigation while MD and Msi Spring were the best.

Furthermore, SD for EC and pH were 4958.3 and 0.60 and ranged from (209 to 21,230) μScm^{-1} and (6.30 to 9.30), respectively. These ranges represent major variations and suggest that most samples were slightly alkaline and saline (Table 1, Fig. 4). Figure 2a, b show the spatial distribution of pH and EC values. The highest EC occurs around Kitelela and Mahoma Nyika, (at the NaCl -rich lithological unit, which are in the middle of the catchment) whilst the lowest value was recorded around Msi Spring.

Local water suitability for grape irrigation based on EC, Na^+ , pH, and Cl^- parameters

The results showed that 14.8% of the samples, including HD for the dry season were above the Na^+ PL for grapes irrigation of about 20.01 meq/L. With regard to EC and Cl^- parameters, 26 and 29.6, 63 and 48.1, 22.2 and 18.5%

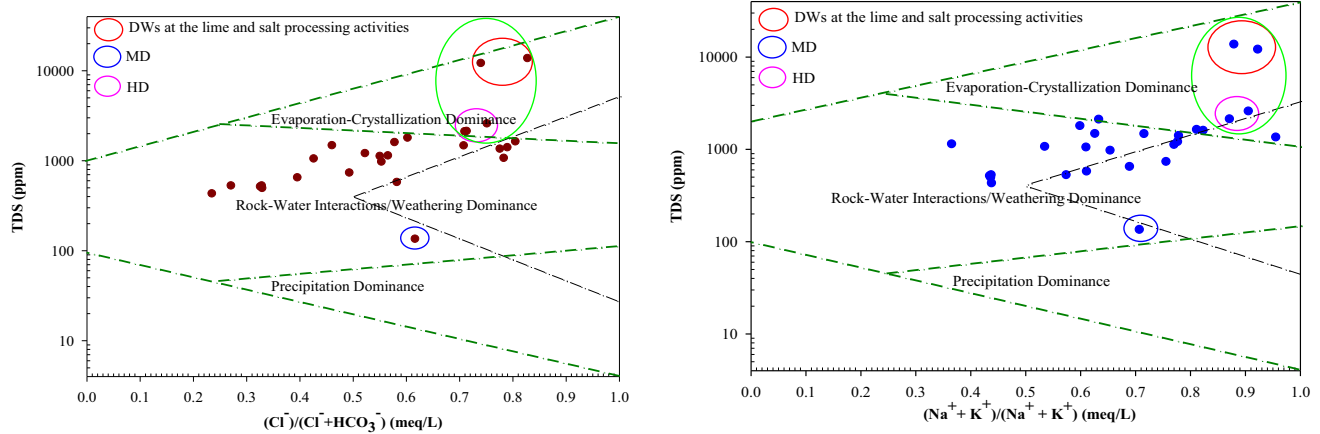


Fig. 4 The major mechanisms controlling the quality of water based on Gibbs (1970)

of the sample would pose none, moderate and severe harm to grape vine, respectively. By using pH parameter, the results show that 96.3% of all study waters could be used for grape vine irrigation (Table 1) while one sample, SN 15 had pH of 9.3 (unsuitable).

Catchment characteristics, human activities and surface hydrology

This survey revealed highly concentrated NaCl-rich lithological materials and limestone rocks that were used to manufacture table salt and lime powder, respectively (Field observations, Plates 1 and 2). A clear layer of NaCl-rich materials was located about 5–7 km upstream of the dam and along the prominent floodplain of LKR that drains to HD via LKR (Fig. 1, Plate 1). These activities heavily use fuelwood, which has resulted to severe catchment destruction (Plates 1 and 2). It is worth noting that these materials are not mapped in the existing geological map (Fig. 1).

This study revealed that the catchment had undergone severe destruction/degradation, due to deforestation, uncontrolled SDWs developments and runoff diversions to SDWs and crop field. There are severe fuelwood shortages due to unregulated vegetation clearances for charcoal/fuelwood (Plates 1 and 2). ‘Getting enough fuelwood at my age is the most challenging part of the salt making process. Recently, we are mostly using crop residues, animal droppings, leaves and grasses or buy fuelwood which, reduces profit margins (Personal communication with Bibi Ekisa Nkazumwa of Nzuguni suburb).

The SDWs were in the runoff channels of HD and are designed to improve well storage and dilute the high water salinity. Field observations showed that the area often floods during heavy rains and when the water recedes to HD, due to irrigation and ET, the concentrated salty water flows with it to the dam.

Discussions

Major mechanism controlling water quality and hydrochemical facies

The major mechanisms controlling compositions and sources of dissolved ions in local water sources are presented by Gibbs and Chadha plots (Figs. 3b, 4). Gibbs plot shows that most samples span in water–rock dominance while a few trended towards ET with none occurring in the precipitation dominance (Fig. 4). Thus, rock–water interactions leading to weathering of the rock forming minerals (ion exchange and reverse ion exchange) were mostly accountable for the main water chemical species. Hence, local geology plays key roles on hydrogeochemistry and water genesis and suggests interactions between the chemistry of rock and that of percolating waters under the subsurface.

This agrees with previous findings that local groundwater dated thousands of years (URT 2002). However, HD for both seasons plotted in ET dominance and closer to SG from the LSPMA (Fig. 4). The co-existence of HD with SG samples suggests common hydrogeochemical mechanisms and is crucial in understanding the dam’s water quality. Worth stressing, although HD plotted in the ET dominance, it does not necessarily imply its hydrochemistry is due to high dam surface ET. The SG from the NaCl-rich lithological unit was equally influenced by ET suggesting that the water type was not exclusive to HD (Fig. 4). The dam also receives water from the groundwater (Shindo 1991) hence its chemistry is also likely influenced by the aquifer properties with the SG at the LSPMA, which was ET dominated, probably making a crucial contribution.

The trending of some samples towards ET was expected for such a semi-arid area with high ET (about 2000 mm/yr.) against low rains (about 550 mm/yr.). Thus, local dryness, semi-aridity and high ET play vital roles in the

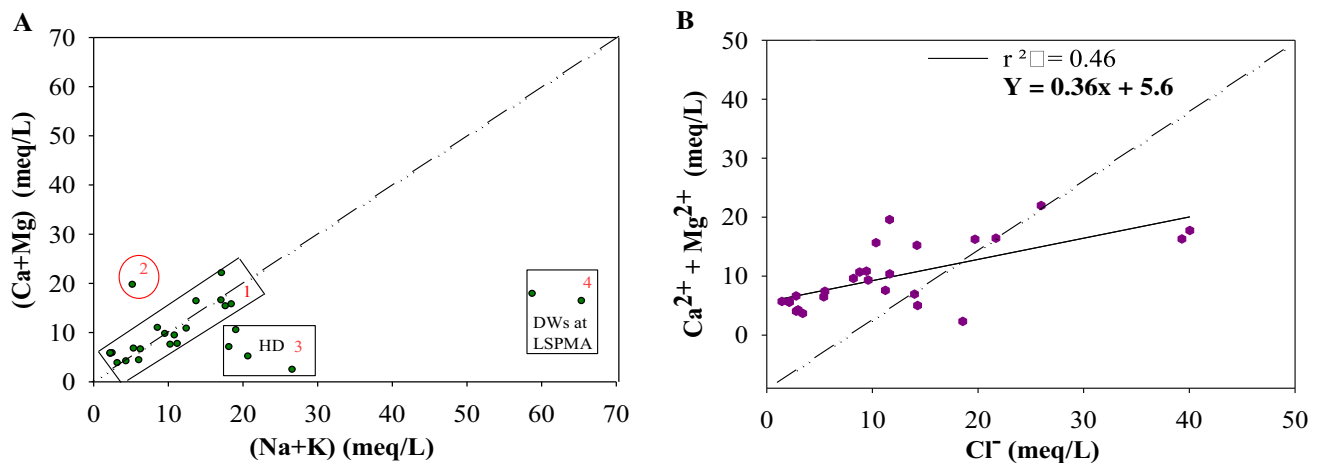


Fig. 5 **a** A plot of $(\text{Na}^+ + \text{K}^+)$ vs. $(\text{Ca}^{2+} + \text{Mg}^{2+})$, **b** A Plot of $(\text{Na}^+ + \text{K}^+)$ vs. Cl^-

local water chemistry and its management. A study in India also showed that ET led to increased chemical budget of groundwater (Kumar 2013).

Unlike HD, MD occurred in rock dominance and closer to precipitation (Fig. 4). These findings offer a sharp contrast between the two dams sitting in similar environmental settings and further challenges the notion that high HD surface ET was the only/main factor for its high salinity as suggested by Shindo (1991).

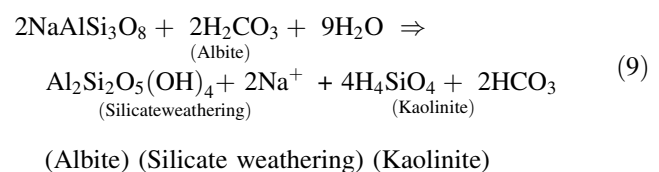
To further explore water chemistry mechanisms, a modification of piper plot after Chadha (1999) showed that most samples plotted in the seawater including samples from HD and SG from the LSPMA area. The next most dominant categories were reverse ion exchange, recharge waters and ion exchange in that order (Fig. 3b). However, Dodoma lies in the inland of the country which rules out the possibility of seawater intrusion. Furthermore, seawater intrusion leads to a plot of $\text{Cl}^-/\text{Anions}$ of >0.8 (Zaidi et al. 2015). Locally, all samples had $\text{Cl}^-/\text{Anions}$ of <0.8 (Fig. 6r) and seawater intrusion was thus excluded. Thus, the water falling in the seawater can only be attributed to excessive ET (Zaidi et al. 2015). Furthermore, the molar Na/Ca ratios of about 52% of samples were >2 (Table 2) indicating silicate minerals reactions and/or occurrence of cation exchange (Prasanna et al. 2011). These results reaffirm that multiple processes control local water chemistry and evolution.

The ratio of Na/Cl vs. EC is also used to assess the roles of ET in water genesis and chemistry. The increasing EC with Na/Cl ratio remaining unchanged indicates that ET is a principal process (Prasanna et al. 2011; Zaidi et al. 2015). Locally, this plot shows a slightly +ve slope suggesting multiple processes other than ET influences genesis of local water sources (Zaidi et al. 2015). To confirm this, Gibbs plot showed that most samples occurred in rock dominance while about 5 trended towards ET (Fig. 4).

Furthermore, in a situation dominated by ET, with no mineral species precipitation, a plot of Na/Cl results in a horizontal line (Fig. 7a). Otherwise, a ratio of >1 is expected when Na^+ is released from silicate weathering (Prasanna et al. 2011). Locally, most samples had a ratio of >1 (Fig. 7a). Higher Na^+ and Cl^- correlation (Fig. 7a) suggests dissolution of chloride salts and/or ET concentration (Jalali 2007). However, the correlation between Cl^- and TA was 0.93 suggesting that most of the alkali ions were balanced by Cl^- (Fig. 6r, Zaidi et al. 2015).

Higher conc. of Cl^- (aver. of 11.83 meq/L, Table 1) suggests longer water–rock silicate weathering interactions (Jalali 2007). This is supported by a poor correlation between HCO_3^- and Cl^- ($r = 0.32$, Fig. 5O) which suggest dissimilar processes were responsible for occurrence of these two ions (Kaka et al. 2011). In contrast, the graph of $(\text{Na}^+ + \text{K}^+)$ vs. $(\text{Cl}^- + \text{SO}_4^{2-})$ resulted into $r^2 = 90$. A strong +ve correlation (i.e. increase in alkalis $(\text{Na}^+ + \text{K}^+)$ corresponded with an increase in $(\text{Cl}^- + \text{SO}_4^{2-})$) implies a common source of these ions (Zaidi et al. 2015; Prasanna et al. 2011).

Figure 7a shows that most samples scatter below the equiline which suggests that reverse ion exchange/halite dissolution were the major mechanisms controlling the bulk of local water chemistry. These include samples from HD and SG from the LSPMA area (Fig. 5a). Furthermore, for most samples, Na^+/Cl^- was >1 suggesting that silicate dissolution/weathering was the most likely source for Na^+ according to Eq. 9 while samples having $\text{Na}^+/\text{Cl}^- < 1$ could be from other mechanisms, including pollution (Kumar 2013; Prasanna et al. 2011).



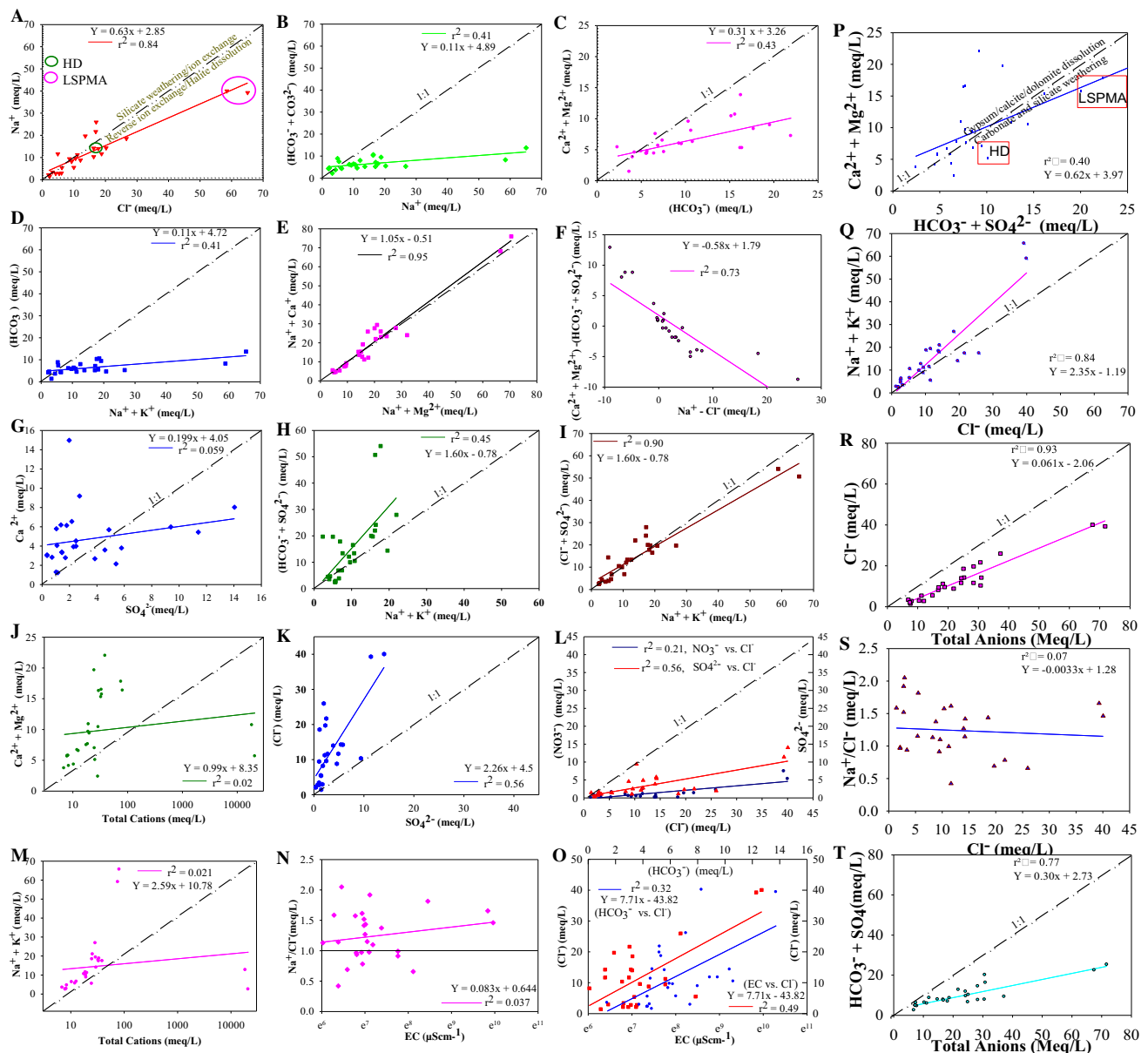


Fig. 6 Ionic relationships and ratios of the major local anions and cations in water sources from Hombolo dam catchment, Tanzania. The dash-dotted lines represent equilines (1:1)

However, Fig. 4 showed that HD together with SG from LSPMA area were in ET dominance which suggest that the water- rock influenced water (Fig. 7a) was further concentrated by ET (Fig. 4). Co-influencing of these processes has widely been reported (Prasanna et al. 2011). The fact that most samples plotted below the equiline ($Na=Cl$) is an additional confirmation that the excess Cl comes from additional geochemical processes, mostly ET and pollution (Prasanna et al. 2011).

Figure 6 l shows simultaneous co-existence of NO_3^- and Cl^- ($r^2 = 21$) which reflects pollution (Neill et al. 2004). The area is highly cultivated and LSPMA take

place with no sanitation facilities in place (Plates 1 and 2). Free range livestock feeding practices involving huge trails of herds were also common. The main wastewater stabilization ponds (WWSP) in the near upstream of the catchment have a very low efficiency (Mkude and Saria 2014). The effluent of the poorly treated wastewater passes through the study area before entering HD. Thus, part of these ions could be related to pollution pre dominantly from sewage and agricultural activities. Notice that the highest NO_3^- was recorded in one of the two SDWs within the LSPMA (SN 27 in Table 1).

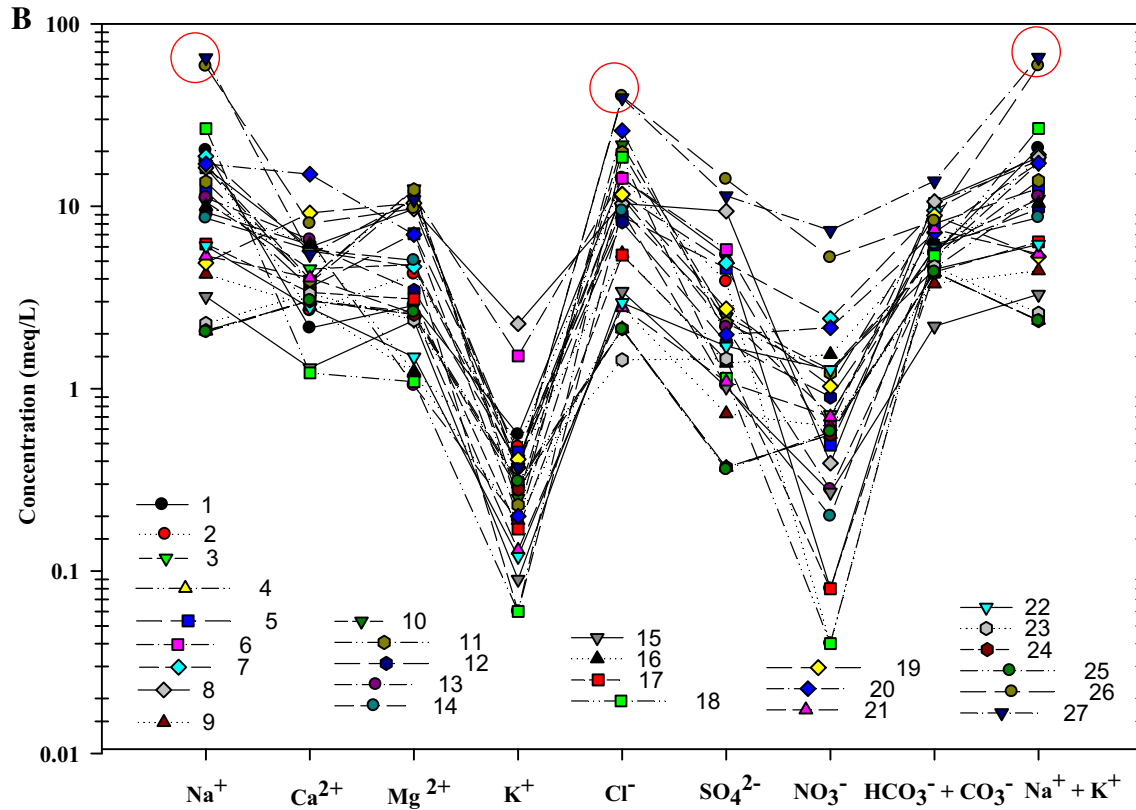
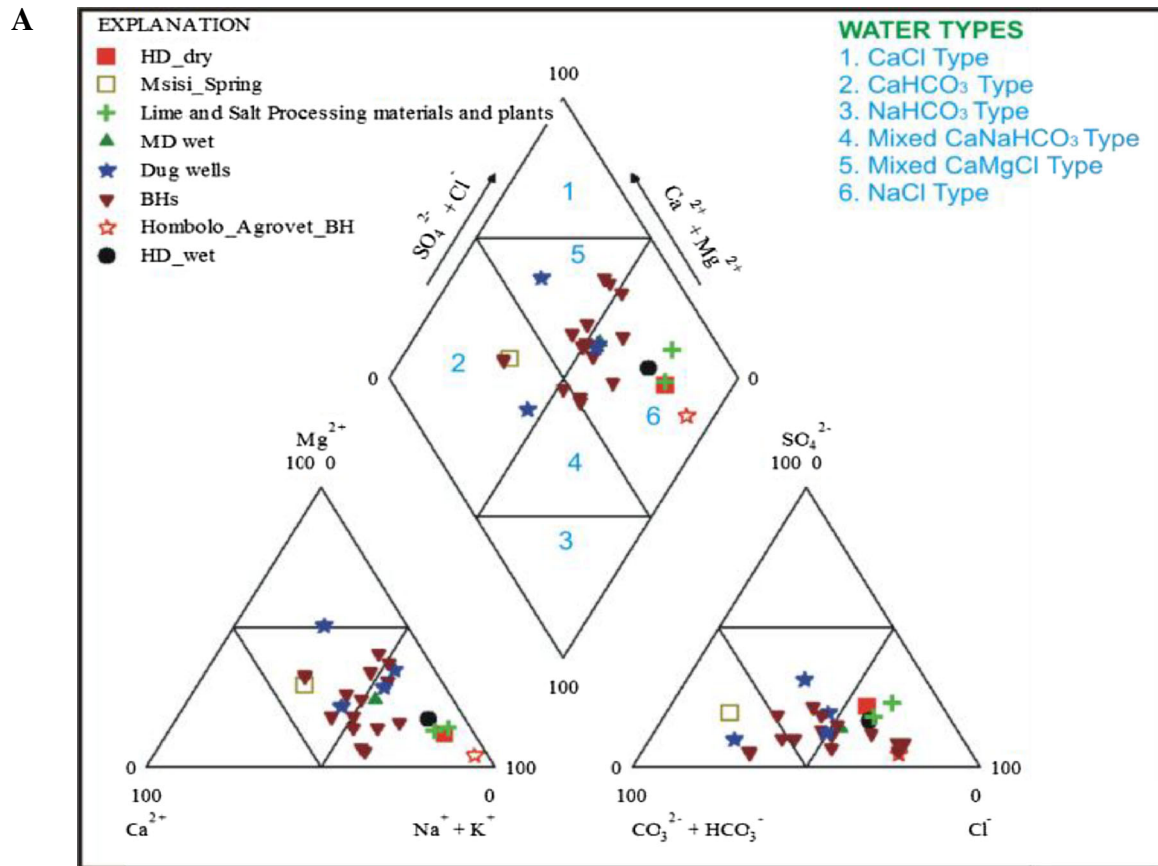


Fig. 7 a Piper Trilinear Diagram, b Schoeller Diagram from the study site

Although Na^+ was the most dominant cation, Fig. 3a shows that alkaline earth metals (AEMs) dominated over alkali metals (AMs) and that four classes could be classified. The first class consisted of 74% and exhibits equal amounts of AEMs and AMs. The second showed a slight abundance of AEMs more than AMs (4% of samples). The third consists of a slight abundance of AMs making 15% and included samples from HD. The fourth was marked by an excess of AEM constituted 7% and included samples from the LSPMA area. The dominance of AEMs over AMs can be attributed to base ion exchange where Groundwater- Na^+ is replaced with Ca^{2+} and Mg^{2+} from rock/soil matrices (Fisher and Mullican 1997; Prasanna et al. 2011; Zaidi et al. 2015).

A plot of $(\text{Ca}^{2+} + \text{Mg}^{2+})$ vs. $(\text{HCO}_3^- + \text{SO}_4^{2-})$ falls close to the equiline when dissolution reactions of calcite, dolomite, and gypsum are dominant in the media. Data falling along the line indicate that the ions mainly originated from weathering of carbonates and sulfate minerals (Prasanna et al. 2011). Locally, about half of the samples plot above the equiline (Fig. 6p.) suggesting that carbonate weathering is the prime source of Ca^{2+} ions. Samples with values >1 , suggest an excess of $(\text{Ca}^{2+} + \text{Mg}^{2+})$ were

derived from a reverse ion exchange (Prasanna et al. 2011). Thus, high $(\text{Ca}^{2+} + \text{Mg}^{2+})$ over $(\text{SO}_4^{2-} + \text{HCO}_3^-)$ strongly suggests a reverse ion exchange (Zaidi et al. 2015). However, the fact that $(\text{Ca}^{2+} + \text{Mg}^{2+})$ increased with increasing salinity (Cl^-) (Fig. 4) suggests more reactions other than reverse-ion exchange were occurring in the clay/weathered layer (Prasanna et al. 2011). Ion exchange reactions tend to shift the points to the right of the 1:1 line due to the excess supplies of $(\text{HCO}_3^- + \text{SO}_4^{2-})$ (Fisher and Mullican 1997). Furthermore, when there is an excess of $(\text{Ca}^{2+} + \text{Mg}^{2+})$ than that of $(\text{HCO}_3^- + \text{SO}_4^{2-})$ ions, often the points tend to shift to the left side of the 1:1 line owing to reverse ion exchange reactions. To confirm ion exchange reactions were not the sole mechanisms behind the observed water chemistry, a graph of $(\text{Ca} + \text{Mg} - \text{SO}_4 - \text{HCO}_3)$ vs. $(\text{Na} - \text{Cl})$ was prepared (Fisher and Mullican 1997). Water exhibiting ion exchange plots along a line with a slope of -1 in $(\text{Ca} + \text{Mg} - \text{SO}_4 - \text{HCO}_3)$ versus $(\text{Na} - \text{Cl})$. Figure 6F shows that the slope was about -0.58 suggesting that although ion exchange was taking place, it was not the only factor (Fisher and Mullican 1997). Thus, samples exhibiting reverse ion exchange could be responsible for the



Plate 1 Representation of major processes involved in salt making plants **a** Removal of top soil in order to collect the inner and more salt concentrated materials at Nzuguni. Notice the whitish salty richer layer just below the top soil. **b** A lady collecting water from a SDW in Kitelela needed to separate salt from the soil. Notice that a young goat fell into the SDW while attempting to drink water and the dark color of the water that is often contaminated by wastes from huge trails of livestock. The water used to process salt is often of unknown quality. **c** Local separation of salt done by vigorously stirring soil materials in

jerry cans, pouring the mixture in perforated jerry cans packed at the base with fresh leaves to prevent soil particles from escaping the cans but allow filtrates to filter through. **d** Boiling the filtrates in big pots that expose a large surface area for freshwater to evaporate and leave the salt in the pots. The source of energy could be anything from animal droppings to twigs and poles collected from heavily deforested vegetation. Notice the small size of fuelwood suggesting deforestation of landscape **e** Removing the salt from the pot after boiling for some 3 h **f** Drying salt ready for packaging

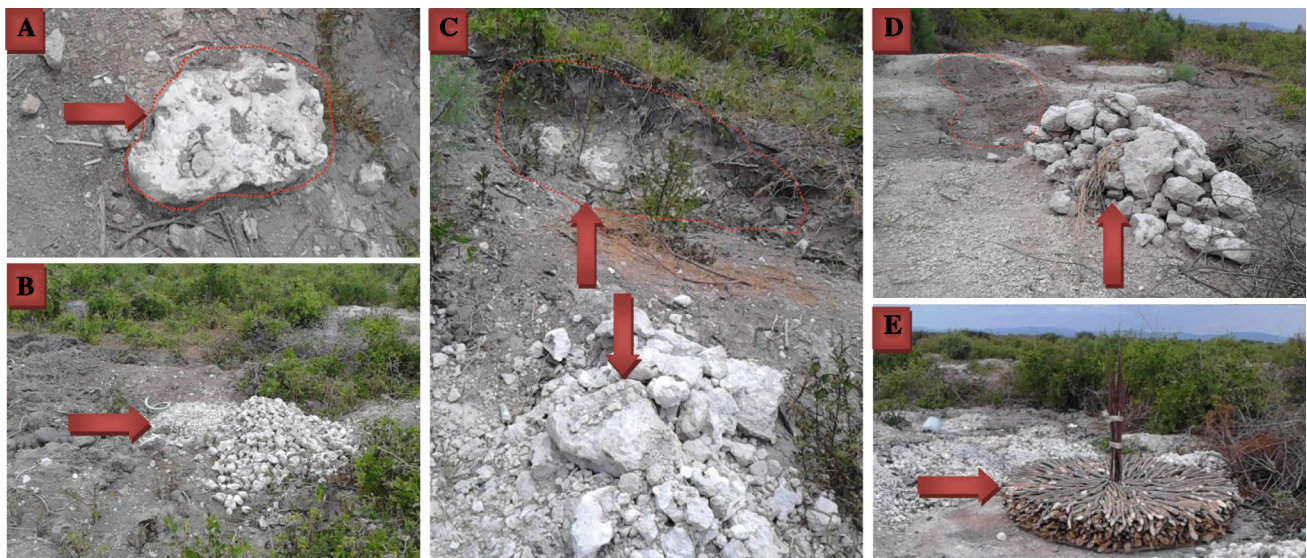
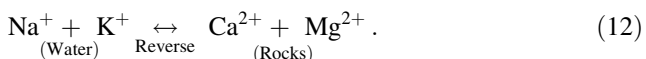
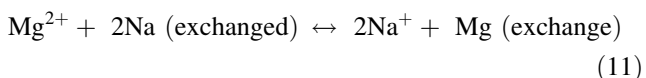
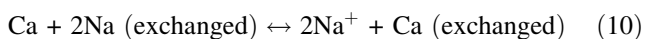


Plate 2 Representation of major processes involved in making lime powder. **a** Partially exposed limestone rock at Kitelela. These are usually dug to expose the entire rock mass, especially when viewed as of a good quality. **b** Limestone rock is broken into smaller pieces to increase surface area for burning and breakage. **c** A scar from where a rock was removed, the lower part of the picture shows initial stages of breaking the rock into smaller pieces. **d** A pile of rocks dug/collected from the catchment. **e** An open furnace made of fire-wood. Notice the small sized fire-woods, indicating that large trees in the landscape had been cleared. Improved designs of closed furnaces like the ones

used to burn/process mud bricks could save energy as a large part of the heat energy is often lost in the open furnace. Such a furnace is often built into several layers by putting fuel wood as shown in D followed by rock pieces in an alternating manner. The process is done in rings and when done the furnace is set on fire for the limestone pieces to burn. When the rocks have burned for some 3–4 h, water is applied on top of the furnace to break the rocks into powder form. Notice the loose vegetation density in the backgrounds of all pictures suggesting the landscape were severely deforested for fuel wood energy needs

depletion of Na^+ with respect to Ca^{2+} and Mg^{2+} (Zaidi et al. 2015).

Moreover, hydrogeochemical properties and changes along its flow paths can be traced via CAI 1 and CAI 2 indices. Locally, 22.2 and 77.8% of the samples had +ve and -ve CAI 1 and CAI 2 depicting reverse ion exchange and ion exchanges, respectively (Fig. 3a, b). The +ve ratio shows the likelihood of exchange of Ca–Mg (water) by Na–K (mineral) (Eq. 12, Kumar 2013). This further suggests that both reverse ion and ion exchange are common processes. However, local water table, particularly from shallow aquifers are often very high (Shindo 1991) which suggest a likelihood of active exchange of soil/clay ($\text{Ca}^{2+} + \text{Mg}^{2+}$) with Na^+ in groundwater (Eq. 10, 11 and 12); where (exchanged) represents cations exchanged on water or soil/clay.



The graph of ($\text{Ca}^{2+} + \text{Mg}^{2+}$) vs. HCO_3^- explains the origin of Ca^{2+} and Mg^{2+} in groundwater (Prasanna et al. 2011). When Mg^{2+} and Ca^{2+} only originate from the

dissolution of carbonates in the aquifer materials and/or weathering of silicate minerals (amphiboles and pyroxenes), this ratio is around 0.5 whereas the ratio of <0.5 suggests depletion of Mg^{2+} and Ca^{2+} relative to HCO_3^- due to ion exchange and/or enrichment of HCO_3^- (Sami 1992). Locally, 88.9% of samples, including from HD and the LSPMA had ratios of >1 (Fig. 6c). This indicates that carbonate weathering was a key source of high HCO_3^- . This is supported by the presence of numerous limestone rocks (Plate 2). However, the few values with low ratio (<0.5) suggest HCO_3^- enrichment and/or ($\text{Ca}^{2+} + \text{Mg}^{2+}$) depletion by cation exchange (Fig. 6c). Nonetheless, since HCO_3^- does not form carbonic acid (H_2CO_3) under mostly alkaline conditions locally (Table 1), the high ratio was unlikely to be due to HCO_3^- depletion (Shindo 1991; Prasanna et al. 2011). However, the high ratio could indicate that an excess of alkalinity was balanced by the alkalis metals (Na^+ and K^+) (Prasanna et al. 2011) or were due to reverse ion exchange (Fig. 3a). Further, the ratio of <1 may also indicate existence of fresh recharge (Zaidi et al. 2015). Locally, however, most samples had a high ratio (average 1.59) suggesting limited fresh recharge which is also supported by the low rains (550 mm/yr.) and higher ET (about 2000 mm/yr.) all of which affect net recharge (Shemsanga et al. 2015). Previous studies also indicated

that the area had low recharge (Shindo 1991, Rwebugisa 2008).

The graph of $(Ca^{2+} + Mg^{2+})$ vs. TC, showed that 85.2% of samples were above the equiline indicating a dominant contribution of $Ca^{2+} + Mg^{2+}$ in the overall water chemistry (Fig. 6j). This is in line with the results shown in Table 1 that Ca^{2+} and Mg^{2+} were the second and third most dominant cations, respectively, and is supported by presence of limestone rocks (Plate 2).

The plots of $(Mg^{2+} + Ca^{2+})/salinity (Cl^-)$ and Na/Cl offer insights on interactions between Ca^{2+} , Mg^{2+} , and Na^+ in the groundwater system (Sami 1992). When all Cl^- is derived from halite dissolution, the increase in $(Ca^{2+} + Mg^{2+})$ corresponds to increasing (Cl^-) . The plot of $(Mg^{2+} + Ca^{2+})$ vs. (Cl^-) showed that $(Ca^{2+} + Mg^{2+})$ increased with the increasing Cl^- (Fig. 5b) whilst (Na/Cl) ratio slightly decreased with increasing Cl^- (Fig. 6s). Thus, the dissolved groundwater Na^+ (derived from halite dissolution) is exchanged by $(Mg^{2+} + Ca^{2+})$ in the clays within the aquifer/soil matrix (Rajmohan and Elango 2004).

Based on the diamond-shaped fields of the piper plot, local water samples were grouped into four major facies. The position of the chemical ions offers vital evidence on chemical history of waters and signifies the catchment characteristics. The dominant water types were $Na-Cl$ (primary salinity) followed by mixed $CaMgCl$. Msisi Spring and one sample each from BHs and SDWs (all collected from around Msisi and Mkoyo suburbs) diverted leftwards towards the $CaHCO_3$ type (Fig. 6). The fact that HD retained the same water type in both seasons suggests poor dilution, i.e., no major impacts shown by rainfall (Fig. 6). No samples represented $CaCl$ and $NaHCO_3$ water while one sample diverted towards $CaNaHCO_3$ type. Except for samples collected from the LSPMA, most SDWs occur together with BHs indicating similar mechanisms (Fig. 6). HD and water collected from the LSPMA area co-existed in the plots which suggests similar water types and evolution mechanisms (Fig. 6). This similarity could partly be due to active HD–groundwater interactions (Shindo 1991) and reinforces the notion that HD salinity is not fully due to ET.

In terms of anions, Cl^- had the most abundance followed HCO_3^- while SO_4^{2-} was completely absent (Fig. 6). The piper diagram shows that the cation and anion triangles were well-defined with most waters falling within the influence of $(Na^+ + K^+)$ and Cl^- , respectively (Fig. 6). While $Na-Cl$ mostly indicated ET controlled water, mixed $CaMgCl$ represent evolved water, meteoric signature of which was lost by rock–water interactions (Zaidi et al. 2015). Thus, the dominance of $Na-Cl$ type could be influenced by the LSPMA (Plate 1) and the extreme semi-arid ET (Shemsanga et al. 2015). This is

supported by the fact that water samples collected away from the LSPMA area, and which acted as control like Msisi Spring had $CaHCO_3$ (carbonate hardness type), probably due to rich limestone outcrops locally (Plate 2). Similar findings were reached by (Rwebugisa 2008).

Although MD water was categorized as $Na-Cl$, it plotted to the extreme left of HD and closer to most BHs and SDWs collected away from the LSPMA area. Generally, MD was the most mineral free sample with its water chemistry closely approaching the precipitation dominance (Fig. 6). Properties of these less mineralised samples could reflect the normal water chemistry representing the average catchment characteristics other than the LSPMA area. Previous studies got comparable results for the rest of the catchment other than the LSPMA area (Nkotagu 1996; Rwebugisa 2008).

Water usability based on hydrogeochemical parameters, national and WHO guidelines

An excess conc. of some parameters causes health hazards (WHO 2008). Locally, the abundances of major dissolved ions were highly variable (Table 1) which partly suggests multiple hydrochemical sources and processes influence local water quality (Prasanna et al. 2011).

Water pH measures the intensity of acidity and/or alkalinity in which all chemical and biological processes depend on (Islam et al. 2016). The results showed that most samples were within the MPLs of both national and WHO standards (96.43%) while SN 18 had a high/unsuitable pH of 9.3. However, 81% of the samples were above the natural water pH range of (4.5–7.0) and suggest possible water quality adjustments (Kaka et al. 2011). Figure 6b shows the spatial distribution of pH where HD and wells around the LSPMA had higher values while samples further away had pH, approaching local rain pH of 6 (Shindo 1991; Nkotagu 1996). Thus, the low pH in SN 3, 4 and 24 could indicate that those areas receive frequent rain recharge which locally has low pH or are frequently pumped and hence less conc. of dissolved minerals (Napacho and Manyele 2010). The latter was especially exemplified and confirmed by SN 4 which was found to be almost continuously abstracted for irrigation purposes (Filed observations, FGDs). However, organic matter decomposition could also lead to low acidity (Ibid). Assessment of pH value is crucial as low pH could indicate presence of toxic metals while higher pH leads to slippery feel and soda taste in water. Low pH water is also soft and corrosive and can leach metals from pipes and fixtures like copper, iron, lead, manganese and zinc (Napacho and Manyele 2010; Kaka et al. 2011). Low pH waters also cause esthetic problems and bitter, metallic to sour taste and can also stain laundry and sanitary sinks (Kaka et al. 2011).

TDS value is vital for assessing longevity of water in the catchment with high values indicating long residence time (Elisante and Muzuka 2016). The results suggest that most samples had high TDS values which shows longer residence time (Table 1). 88.9 and 18.5% of the samples were above the recommended WHO and national PL, respectively, and highlights the big differences between the two guidelines. Similar findings were reported by Rwebugisa (2008) and Shindo (1991). Generally, areas of low TDS denote recharge zones and are referred to as young waters, whilst those with high TDS tend to be discharge sites and referred to as old waters (Kaka et al. 2011; Elisante and Muzuka 2017).

High conc. of Mg^{2+} and Ca^{2+} causes of higher water alkalinity, high TH, authentic problems and bitter water taste (Napacho and Manyele 2010). Based on TH classification, 66.7, 29.3, and 3.7% of the local samples were grouped as very hard, hard and moderately hard, respectively (Table 1). This shows that high TH is a serious problem locally. Previous studies reached similar findings (Shindo 1991; Nkotagu 1996; Rwebugisa 2008). Generally, hard water requires more soap, i.e., extra resources that could be used for other uses in such an area where average households earn <1USD/day (URT 2013). High TH water has also been linked with cardiac disorders (Kaka et al. 2011) hence must be handled with care in such area with ill functioning health services (Ibid).

The results show that 3.7% of the samples were above Ca^{2+} PL for both WHO and national standards while 3.7 and 22.2% of the samples were above the WHO and national Mg^{2+} guidelines, respectively (Table 1). Consuming water with high Ca^{2+} affects absorption of other vital minerals and is linked to kidney stones (Eletta et al. 2010). While high Mg^{2+} in water leads to undesirable taste and a laxative effect (Ibid), low Mg^{2+} causes increased mortality from cardiovascular diseases and is associated with higher risk of motor neuronal diseases, pregnancy disorders, and preeclampsia. Low Ca^{2+} in water is linked to risk of neurodegenerative diseases in children, pre-term and low weight births. Shortage of both Ca^{2+} and Mg^{2+} in water is associated with some forms of cancer (Napacho and Manyele 2010). The common source of Ca^{2+} and Mg^{2+} in groundwater is erosion of limestone and dolomite rocks/minerals, such as calcite and magnetite (Marques et al. 2014). While the catchment is rich in limestone rocks, the relatively low Ca^{2+} and Mg^{2+} locally could be due to ordinary softening by cation exchange by Na^+ rich materials displacing Ca^{2+} and Mg^{2+} in solution (Nkotagu 1996).

Table 1 shows that 48.1 and 11.1% of the samples were above WHO and national PL for K^+ , respectively. These results suggest that K^+ is a serious problem locally. In natural environment, however, K^+ often occurs in low

levels owing to its tendency to be fixed by clay minerals and its involvement in the formation of secondary minerals (Pazand et al. 2012). The origin of K^+ in natural freshwater is weathering of rocks but its conc. often increase in the polluted water due to disposal of wastewater (Pazand et al. 2012). Otherwise, the commonest K^+ minerals are orthoclase, feldspar, microcline, leucite, biotite in local granites of the area (Nkotagu 1996).

The results show that 72.7 and 11.1% of samples were above the WHO and national PL for Na^+ , respectively (Table 1). Thus, Na^+ was mostly higher than the WHO PL of 8.7 meq/L. High conc. of Na^+ from most local water samples could be due to long rock–water interactions, high ET, dissolution of halite minerals and weathering of plagioclase feldspar (Pazand et al. 2012). The weathering of plagioclase feldspars (albite) follows Eq. 9. While, low Na^+ is recommended for hypertension, congenital heart and kidney patients (Napacho and Manyele 2010), local values represent a threat to these vulnerable individuals.

Table 1 shows that 37 and 30% of the samples were above NO_3^- PL for WHO and national standards, respectively. The highest NO_3^- values were found at the LSPMA area which had no sanitation facilities (FGDs and Filed observations). However, the high value of NO_3^- locally could also be due to oxidation of ammonium formed from sewerage passing through the unsaturated zone, where sufficient O_2 diffuses from the atmosphere, causing oxidation of the reduced wastewater (Stumm and Morgan 2012). In contrast, low NO_3^- in some samples suggests little pollution and/or the underlying natural geology in those areas did not have the anion (Kaka et al. 2011). Since NO_3^- is extremely water soluble, it is very difficult to manage it once in the catchment (Elisante and Muzuka 2017, 2016). Thus, the best way to deal with NO_3^- in groundwater system is to avoid catchment contamination. Drinking water with high NO_3^- leads to methemoglobinemia (blue baby syndrome) (Rebelo et al. 2015). High NO_3^- is associated with risk of bladder cancer, hypertension, increased infant mortality, central nervous system, birth defects, diabetes, spontaneous abortions, respiratory tract infections, and changes to the immune system (Majumdar and Gupta 2000). Thus, water samples exceeding the PL for domestic and livestock uses should be avoided.

This survey also revealed that 22.2 and 11.1% of the samples were above the WHO and national PL for Cl^- , respectively (Table 1). An excess of Cl^- renders a salty taste in water and a laxative effect to people not accustomed to it (WHO 2008). Apart from geological reasons, high Cl^- may be due to poor sanitation (Mkude and Saria 2014) and high ET (Shindo 1991). Based on Soltan, (1998) system of classifying water based on Cl , SO_4^{2-} , HCO_3 cons., water sources can be grouped as normal chloride

(<15 meq/L), normal sulfate (<6 meq/L), and normal bicarbonate (2–7 meq/L) types. By following this method, 78, 89 and 74% of samples could be classified as normal chloride, normal sulfate and normal bicarbonate water types, respectively (Table 1).

While all samples were within the national SO_4^{2-} PL, 11.1% were above the PL for WHO (Table 1). The low SO_4^{2-} locally suggest limited supply of SO_4^{2-} minerals in the underlying geology (Kaka et al. 2011). Drinking high SO_4^{2-} leads to cathartic, respiratory, diarrhea and dehydration (Subba Rao 1993). Further, high SO_4^{2-} , lead to an offensive taste in drinking water, interferes with the efficiency of chlorination in water treatments and increase water corrosiveness (WHO 2008).

Analysis of surface hydrology and recent catchment/water sources salinity increases

A clear layer of highly concentrated NaCl-rich lithological unit and limestone rocks were located in the LKR floodplain that drains into HD (Plates 1, 2). The formation of such layers of salt from near the ET surfaces in dry areas is common, especially when restricted by limited freshwater exchange (Karanth 1994). The study area is classified as semiarid and has very high ET, about 2000 mm/yr, almost 4 folds the amount of rainfall, (550 mm/yr.) (Shemsanga et al. 2015). To account for high conc. of the NaCl-materials, craftsmen in Nzuguni, Kitelela and Mahoma Nyika were actively producing table salt from these materials. The salt making process is done by (1) collecting and soaking salty soils in jerry cans, (2) vigorously stirring and filtering the mixture to get salt concentrated water, (3) boiling the filtrates to make salts (Plate 1). Between 500 and 1000 craftsmen work in about 60 SPPs (depending on seasonality with the peak of the dry season being more productive/preferred). The fact that the LSPMA occur in the direct drainage to HD, suggests a possible conduit of the salinity to the dam (Fig. 1).

The extent of the soil salinity (conc.) can be appreciated from the fact that 50L of filtrated soil water produce about 20 kg of table salt (Filed observations and FGDs). According to mama Ombeni of Kitelela, the net salt production is slightly higher in the dry season than in the wet season by a difference of ~3 kg/every 50L of filtrated water. FGDs revealed that one individual produces about 50 kg of salt, which sells for 0.25\$/kg, an average of 12.5\$ in one day. This is considered a good income locally where per capita income is <1USD/day (URT 2013). Thus, the LSPMA are likely to continue for a long time, and in order for them to be sustainable, must be regulated. This is especially the case owing to recurring poor harvests in recent years caused by repeated shortages of rainfall (FGDs). These findings are supported by recent climate

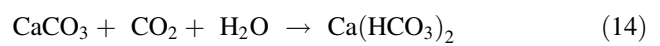
trends that show worsening longevity and timing of growing seasons (Shemsanga et al. 2015).

The same area had numerous limestone rocks which were used to make lime powder and employing about 200 craftsmen (Plate 2). Macheyeke et al. (2008) also report that calcium rich (calcrete cement) materials that fill high-angle fractures were formed close to the subsurface with alternating periods of flooding and drying, in a geographically low-relief environment within the Bahi depression in which the HD lies. Limestone, a water soluble sedimentary rock, is excavated, broken into small pieces, burned on high heat open furnaces and plenty of water is then added to the heated rocks to break them into powder. These processes increase the surface areas for water dissolution and reaction with minerals and make it easier for runoff to dissolve the salts and carry them downstream to HD. The burning of limestone rocks follows thermal Eq. 13:



(Ghalibaf and Moussavi 2014).

The presence of limestone locally is important since it is the major cause of water hardness (Napacho and Manyele 2010). CaCO_3 reacts with H_2O that is saturated with CO_2 to form soluble $\text{Ca}(\text{HCO}_3)_2$ according to the Eq. 14: Previous studies also indicated that high water hardness was a serious problem (Shindo 1991; Nkotagu 1996; Rwebugisa 2008).



(Jiang et al. 2012).

Further, limestone dissolves to release $\text{Ca}^{2+} + \text{CO}_3^{2-}$ ions (Eq. 15): It is, therefore, not surprising that HD and most local water sources had high proportions of HCO_3^- anions/species (Table 1).



(Jiang et al. 2012).

During heavy rains and mostly when the dam is full, HD water floods back to the catchment including to the LSPMA (i.e. NaCl-rich lithological unit). It is likely that part of salts potentially dissolves, percolate and/or are gradually carried down to HD when its water progressively recedes from ET and/or abstractions (Field observations). The micro-recharge is thought to be responsible for high salinity in the SG at the NaCl-rich lithological unit (Fig. 1, Table 1). The likelihood of the micro-recharge was supported by the spatial distribution of EC where areas around NaCl-rich lithological unit had high EC while those further away had lower EC values (Fig. 2a). However, HD observes a slight dilution with less mineralised/fresh runoff where EC and pH dropped by about $708 \mu\text{Scm}^{-1}$ and 0.5, respectively (Table 1). As ET continues to remove water from the dam, freshwater molecules escapes, leaving more salt ions hence HD progressively becomes saltier again.

However, even during the wet season, HD retained relatively higher salinity (EC) than recommended values for irrigation and domestic (Table 1). Thus, the idea that HD salinity is fully due to dam surface ET is uncertain and highly disputed. It was clear that the salts were coming with runoff prior to being concentrated by ET. The chemistry of the SG within the LSPMA was ET dominated and explains why HD water was in the same category. Thus, the local groundwater quality is associated with high ET and HD—groundwater interactions (Shemsanga et al. 2015).

At this point, the claim that HD is recently saltier remains unresolved. However, HIS only started in 1995 and mostly operates during wet and a few months thereafter (when the water is less salty) (Field observations and personal communication with Mr. Ntuza). HIS brought in a new factor that withdraws water and largely reduce HD volume when the salinity is relatively low and could account to recently increased salinity. Similarly, from 1921 to 2014, there was a decline of about 54 mm of annual rains. Shemsanga et al. (2015) showed that except for 6 mm of rains recorded in 2009, there had been no May rains from 1991 to 2014 whereas May used to contribute significantly to HD in previous years. This trend means less freshwater in the catchment and poor HD dilution hence increased salinity recently. Furthermore, Field observations revealed several runoff diversions upstream of HD. These diversions trap runoff flowing to HD (Shemsanga et al. 2015). The fact that the diversions were mostly from salt-free areas, means less freshwater to HD and more solutes per unit dam volume. It is also argued that HD level has reduced recently owing to these diversions (FGDS and Personal communication with Mr. Ntuza).

A recent study also indicated a warming trend where between 1964 and 2014, T_{\min} increased for about 1.13 °C (Shemsanga et al. 2015). The warming in turn triggers more ET, which implies additional water losses and concentrating catchment's salinity, including in HD. This could further explain recent HD salinity increases corresponding to increasing PET/T and decreasing rainfall. The local warming is crucial as, T and ET are highest in wet months hence more water losses (Ibid). This is especially worrisome since wind, another factor driving ET, is also highest during wet months and local rains are often interrupted by strong sunhours (Ibid). These characteristics explain the lack of permanent water bodies, except for HD which partly receives groundwater (Shindo 1991). These factors are argued to have reduced HD level to the average of 4.5 m compared to the original design of 9 m (personal communication with Mr. Ntuza). However, HD sediments account for large part of the dam volume which exposes enormous dam volume to direct ET forces/losses.

The gradual increase in groundwater salinity is also not fully resolved at this time. However, poor recharge due to low rains and/or high ET leads to unlocking of deeper groundwaters that tend to be mineral richer leading to pumpage of deeper waters in contact with rock matrices (Tularam and Krishna 2009). Since local recharge highly depends on local rainfall; the decreasing rains have the potential of affecting groundwater quality (Taylor et al. 2012, Shemsanga et al. 2015). Decreasing rains/recharge is supported by local groundwater level trends (Shemsanga et al. 2015). Similarly, there are active HD—aquifer interactions; thus the gradual groundwater salinity increase could originate from HD salinity increases (Shemsanga et al. in preparation). This is supported by the fact that most wells close to HD had saltier waters while wells far away, and which acted as controls, had low salinity (Shemsanga et al. in preparation). It is worth mentioning that HD is a prominent discharge site for Makutupora aquifer, the main aquifers supplying water in DM. The possibility of these two water masses mingling requires careful understanding of these water components and further shows the need for holistically managing water resources, i.e., integrated water resources management (URT 2009).

Comparison of water suitability/compositions between HD and MD

The origin and mechanisms of high salinity in HD can be appreciated when compared to a nearby MD (only about 38 km apart). MD, also an earthen dam, is located within climatologically similar area and thus faces more-or-less similar climatological forces, including high ET rates. However, MD had very low EC ($209 \mu\text{Scm}^{-1}$) and pH (6.6) compared to ($3298\text{--}4006 \mu\text{Scm}^{-1}$) and (8–8.5) of HD (Table 1). Similarly, MD had low conc. of all chemical water quality parameters, particularly, Na^+ , HCO_3^- and Cl^- ions that especially characterize HD (Table 1). Most MD hydrogeochemical parameters were closer to local rain values suggested by Shindo (1991). Thus, HD high salinity is not fully a function of high PET rates but rather a physical source entering the dam.

Salt and limestone manufacturing and environmental degradation

Although LSPMA give locals a good income, they use a lot of fuel wood for which case most trees have been cleared (Plate 1 and 2). The extent of land cover clearance is so significant that not enough fuel wood is currently available and locals use cattle droppings, agricultural remains, grasses, leaves, twigs and poles to supply energy in the stoves (Field observations and personal communication

with Bibi Ekisa Nkazumwa, Elis Malambati of Nzuguni suburb). The deforestation has major environmental implications, particularly by triggering erosion and sedimentation of HD and must be managed (Field observations and personal communication with Mr. Ntuza). Severe catchment modification is also linked to higher runoff and reduced infiltration (Valimba 2004) hence may affect groundwater recharge locally. Recent discoveries of natural gas in the southern part of the country (Demierre et al. 2015); can supply needed energy and slow down deforestation rates. Furthermore, instead of using open furnaces to produce lime powder, energy saving stoves and closed furnaces which use less wood like the ones used to make local mud bricks can be adopted (URT 2015). FGD also reveals that there were no sanitation facilities in the LSPMA area and craftsmen mostly relieved themselves in the bushes. This trend is supported by water quality data where the two SDWs from LSPMA area had highest NO_3^- of 5.23 and 7.37, respectively, compared to the catchment average of 1.13 meq/L (Table 1).

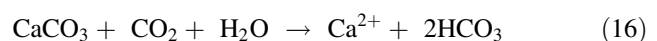
Irrigation water suitability based on existing indices and physico-chemical parameters

Suitability of water for irrigation is a function of its mineral compositions (Table 2). The quality of water can have an adverse impact on soil/crop productivity. An excess of even small amounts of some salts, but when applied for a long time, coupled with poor agricultural practices and drainage, may radically affect soils (Inayathulla and Paul 2013). High conc. of Na^+ in water reduces soil permeability and soil/crop productivity (Ibid). Hence, assessment of Na^+ conc. in irrigation water is crucial and is often represented as %Na. Table 2 shows that, 77.7% of the samples were grouped as excellent (3.7%), good (7.4%) or permissible (59.3%). However, (7.7%) and (3.7%) of samples were classified as doubtful and unsuitable, respectively, and could harm crops/soils. These include HD samples (70.5–78.4%, C4), the two SG samples from the LSPMA (76.3–79.6, C4) and SN 18 (91.8%, C5) and should be avoided for irrigation. Generally, water with Na% of >60% may result in Na accumulation and deterioration of soil structure, infiltration and aeration rates (Inayathulla and Paul 2013). Further analysis of MD water shows that its EC, %Na, SRC, KR values were all acceptable which is a sharp contrast to HD where most of these parameters were unacceptable and/or doubtful (Table 2). The unfitness of HD water for irrigation is distressing as it already employs 468 farmers in the HIS which supports an annual per capita income of 4000\$ (1,171,200\$ of annual local income) (Personal communication with Mr. Ntuza). Thus, sustainability of HD is vital to the populace.

By using SAR index, nearly all samples were either excellent (S1, 85.2%) or good (S2, 3.7%), i.e., suitable for all crops and soils (Inayathulla and Paul 2013). In contrast, 55.6% of the local samples had KR of >1 hence unfit for irrigation (Inayathulla and Paul 2013). These include HD samples for both seasons and challenge its main objective as a source of irrigation water (Table 2). RSC significantly influences irrigation water pH, EC and SAR values. Water with RSC of >2.5 meq/L leads to salt build up which clogs soil pores and hinders air and water movements (Inayathulla and Paul, 2013). Locally, 96% of the samples were in C1 suggesting fitness for irrigation while 1 sample (SN 18) was unfit with a value of 3.031 meq/L (Table 2). In addition, most samples (88.9%) had –ve RSC notations suggesting even desirable fitness for irrigation. –ve RSC reduces soil Na hazard since adequate Ca^{2+} and Mg^{2+} would be in surplus of what will be precipitated as carbonates hence their exclusion from water (Sadashivaiah et al. 2008).

While these indices are useful, they only give empirical conclusions; and other factors are equally crucial to consider. These include climate, crop salinity tolerance, soil types/permeability, rainfall characteristics and irrigation methods (Mulwa et al. 2013). For instance, different crops have different salt tolerance and farmers may need to adapt to specific crops. Well-drained soils may render unsuitable water to be fairly usable. Locally, however, this is unlikely due to low rains (about 550 mm/yr.) hence poor drainage (Shemsanga et al. 2015). A combination of several indices, however, helps to effectively map unsuitable water sources (Mulwa et al. 2013). Thus, by combining these indices, it can be concluded that HD and SN 18 were among the most problematic and should be handled with care while MD was among the best (Tables 1 and 2).

Although high HCO_3^- is not toxic to plants, conc. of >2 meq/L causes Zinc deficiency in rice (Asiwaju-Bello et al. 2013). Locally, 96.3% of samples, including HD had values of >2 meq/L (Table 1). While rice is currently not a common crop, these results are useful if rice cultivation will be sought in the future. Furthermore, high HCO_3^- in water lead to precipitation of insoluble Mg^{2+} and Ca^{2+} in soils, thus leaving higher abundances of Na i.e. increase in the SAR ratio (Mulwa et al. 2013, Asiwaju-Bello et al. 2013). Locally, natural sources like the reaction of carbonate minerals with carbon dioxide gas (CO_2) and dissolution of CO_2 may be the origin of bicarbonate water based on Eqs. 16 and 17:



While water with CO_3^{2-} of >0.1 meq/L is unsuitable for irrigation (Mulwa et al. 2013), 92.6% of samples had conc.

of <0.1 meq/L while 7.4% had values >0.1 meq/L and were exemplified by HD for the dry season and SN 18 that was problematic with most other indices (Table 2). However, low CO_3^{2-} water is common in low pH environments (Elisante and Muzuka 2016).

While recommended Cl^- in irrigation water is <4.0 meq/L (Ayers and Westcot 1985), 70.4% of local water samples were unsuitable, including from HD (Table 1). This shows that high Cl^- is a serious problem locally and water should be used with care if the soils are to remain productive. According to Stuyfzand, (1989), Cl^- is usually classified as (<0.14) Extremely fresh; (C1), (<0.14) Very fresh; (C2), (0.14–0.85) Fresh (C3), (0.85–4.23), Fresh brackish; (C4), (4.23–8.46) Brackish; (C5), (8.46–28.21) Brackish Salt; 28.21–564.13) (C6), Hypersaline, >564.13 (C7). Locally, 0, 0, 29.6, 14.8, 48.1, 7.4 and 0% of samples were grouped into C1 to C7, respectively. The majority (48.1%) were brackish and unsuitable for most uses (Table 1). The two SG from LSPMA, 7.4%, fell under (C6, Hypersaline). Other samples were either in C4 (14.8%) or C3 (29.6%) (Table 1). Apart from the geological reasons, the relatively higher Cl^- could be attributed to atmospheric deposits, pollution (Kaka et al. 2011). A recent study showed that the municipal WWSP found within the same catchment had very low efficiency (Mkude and Saria 2014).

For water to be suitable for irrigation, pH should range from 6.5 to 8.4 (Naseem et al. 2010). While 92.6% of samples were fit for irrigation, HD for the dry season and SN 15, (Table 1) had high pH values 8.5 and 9.3, respectively, and challenge the objective of HD as a source of irrigation water. During the wet season, pH for HD dropped to 8.0 indicating a slight shift towards slim acceptability owing to limited rain dilution which locally has low pH values (Shindo 1991; Nkotagu 1996).

The extent of dissolved ions locally was indicated by high EC values (Table 1). High EC variations were observed with the SG in the LSPMA having the highest, while MD had the least (Table 1). Generally, the higher the EC/TDS the poorer is the salinity hazard (Islam et al. 2016). High EC in water leads to the inability of plants to compete with ions in the soil (Inayathulla and Paul 2013). In terms of irrigation suitability, 55.6% of the samples presented average (C3) while 37% were under C4 (bad for irrigation). HD had high EC of 3298 and 4006 μScm^{-1} for wet and dry seasons, respectively, which suggests poor rainfall dilution and/or the water feeding the dam was already saline with both seasons having C4 waters (Table 1). Thus, HD water was mostly unsuitable for irrigation and should it be necessary to use its waters, mixing it with runoff may be required to lower its EC to <2250 μScm^{-1} . Note that MD water scored C1 and offers a sharp contrast with HD, C4 (Table 1). Similarly, Msihi spring

had low EC of about 660 $\mu\text{S cm}^{-1}$ which suggest low dissolved ions, possibly as a result of shorter water–rock interactions due to good aquifer permeability and steeper slopes leading to rapid groundwater flow (Elisante and Muzuka 2016). In order to understand HD salinity mechanisms further, a tributary leading to HD from the LSPMA area was periodically measured for EC and high values (C4) of (7880, 7690, 8789 and 8670 μScm^{-1}) were recorded in April, 2014 hence confirming poor rainfall dilution of the salts.

Water with TDS of <1000 ppm is regarded as fresh and doesn't affect the soil solution osmotic pressure (Al-Tabbal and Al-Zboon 2012). Locally, 63% of samples had TDS of >1000 ppm hence useless to irrigators (Table 1). Notice that, although HD is frequently/ideally used for irrigation particularly grapes, its TDS values were >1000 ppm for both dry and wet seasons. This might suggest that the soils irrigated with such waters will be adversely impacted which can also explain poor crop harvests locally, including grape (FGDs, field observations and personal communication with HIS officers, village leaders, and farmers).

Study water and its suitability for grape irrigation

Since HD is mainly used for grape vine irrigation and that grape plants are sensitive to EC, Na^+ and Cl^- (Tanji and Kielen 2002; Stevens et al., 2012); this study specifically assessed HD water suitability for grape irrigation. From EC parameter, HD water for both wet and dry seasons were unsuitable for grape irrigation while based on Cl^- , only HD sample for dry season was unsuitable (Table 1). High salinity affects vine productivity by reducing their capability to take up water and eventually via physical damages when the conc. of ions reach toxic levels (Pitt and Stevens 2012). These results generally challenge the main objective of HIS and can be among the reasons for reported poor grape productivity in Hombolo suburb (FGDs). Similarly, owing to high salinity in many SDWs, farmers and herders in Nzuguni, Kitelela, Mahoma Nyika and Mahoma Makulu diverted runoff into some SDWs so as to dilute the salinity and make them suitable for irrigation, livestock and domestic uses (Field observations and FGDs).

Conclusions and recommendations

The mechanisms controlling water quality in HD and its catchment were undertaken to assess the origin of high dam salinity that was debated for over 20 years but remained unclear. Water–rock interactions and high ET were found to affect the quality of different water sources. A layer of NaCl-rich lithological unit was found about 5–7 km upstream of HD amongst limestone rocks. These materials

were in the floodplain draining to HD. During heavy rains, the area often floods with runoff and water from HD and dissolves the salts/materials. The salty water, then recharges the aquifer and/or flows back to HD. This mechanism was supported by the spatial EC distribution (which showed that the NaCl-rich lithological unit area had highest EC while samples far away, and which acted as controls, like Msisi Spring had the least). The dominant water type was Na–Cl followed by mixed CaMgCl and CaHCO₃ in that order. Specifically, HD and SG close to the dam had Na–Cl type and mostly controlled by ET; which is common for dry and semi-arid climate.

Certain water quality parameters (mostly NO₃[−] and EC) in some water sources were above the MPLs for domestic, livestock and agricultural uses and it was clear that some local people were mostly consuming water of unknown quality. This has the potential of affecting human and livestock health, but also the soils/crops irrigated with the unsuitable waters. Thus, imminent adverse impacts on soils are highly likely as the catchment is also not well drained due to low rains (550 mm/yr.) and has decreasing rain trends and high ET (2000 mm/yr.).

Unlike MD, HD was highly saline with higher pH. However, the dams are only 38 km apart with similar climate, hence if the dam surface ET was the major factor for HD high salinity, comparable trends would be recorded in MD, which wasn't the case. MD had a relatively low pH (6.6) and EC (209 μScm^{−1}) compared to HD (8–8.5) and (3298–4006 μScm^{−1}) respectively. This further suggests that high PET locally is not the primary and only reason for the high salinity in HD.

Recent increased water withdrawal from HD and adverse shifts in climatic trends, i.e., decrease rainfall and warming trends were also associated with worsening catchment's salinity. While, rising ET and warming remove freshwater and leave behind concentrated salts, low rains reduce dam volume, mainly from salt-free mountainous areas which feeds the dam and the catchment.

HD water and some wells in the near vicinity of the LSPMA were among the most problematic in terms of irrigation suitability. Although HD is principally used for grape vine irrigation, multiple parameters assessed, including EC and Cl[−] clearly showed that its water was unsuitable for grape irrigation. %Na, in particular, suggests that HD was unsuitable for irrigation and so was SG from the LSPMA area. Unlike HD, MD was mostly permissible under KR, SAR, RSC and %Na indices. Should HD water be used for irrigation, it should be mixed with fresh runoff to dilute the salinity. Small earthen dams can be built for that purpose. It was further recommended that HD sediments be removed and deposition of new sediments be controlled to enable it to hold more water and reduced water exposed to direct ET. Similarly, drip irrigation and

other water-saving irrigation methods should be adopted. Finally, efforts should be made to educate irrigators on the agrarian system on the harms of continued use of saline waters.

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